

- 25 • Demonstration for real case of run-of-river hydropower in Nepal.

26 **1. INTRODUCTION**

27 There is increasing interest in the development of hydropower as a source of renewable,
28 clean energy able to increase the penetration of other renewables as a result of its ability to store
29 energy and supply reliable baseload. The hydropower opportunities left undeveloped since the
30 1970s are being re-evaluated due to a combination of the increase in global energy demand
31 (population growth coupled with increasing per capita electricity demand) and the urgent need to
32 decrease greenhouse gas emissions (Zarfl, et al., 2015). As a partial solution to the shortfall of
33 renewable energy, hydropower holds great promise: existing hydropower generation capacity is
34 sufficient to supply the electricity needs of one billion people, and only approximately a quarter
35 of its global potential has so far been developed (World Energy Council, 2016).

36 Thirty-six gigawatts (GW) of new hydropower capacity were added worldwide in 2014,
37 and 33 more were added in 2015, bringing the global installed capacity to over 1,200 GW (IHA,
38 2016). Still, particularly vast hydropower resources remain untapped in the Indus, Ganges, and
39 Brahmaputra river basins of south Asia (Rasul, 2014; Ray, et al., 2015). In Africa, likewise,
40 developed hydropower capacity is approximately 14 GW (Cervigni, et al., 2015), or less than 8%
41 of the 1900 GW of hydropower potential (World Bank, 2009). Six hundred and forty five million
42 Africans have no access to electricity (IHA, 2016), but hydropower project development
43 spending has fallen victim to a continent-wide infrastructure funding gap (Foster and Briceño-
44 Garmendia, 2010). In some countries of Latin and South America, much of the economically
45 exploitable hydropower has been developed (e.g., Uruguay, Venezuela), but in other countries
46 the bulk of hydropower potential remains untapped (e.g., Brazil, Costa Rica, Chile, Colombia,
47 Ecuador, Peru) (World Energy Council, 2013). Overall, there remains an estimated 430 GW of
48 unexploited hydropower potential in the Latin American region (IHA, 2016).

49 Concerns slowing the adoption of hydropower worldwide (with the notable exception of
50 China, which now produces approximately a quarter of the world’s hydropower (OECD IEA,
51 2015)) are often linked to doubts about the long-term resilience of hydropower facilities in a
52 changing climate (Mukheibir, 2013; van Vliet, Michelle T H, et al., 2016). Because of the large
53 capital costs required, as well as up-front social costs (e.g., population displacement) and
54 environmental costs (e.g., flooding of critical habitat), potential regrets associated with
55 investments in hydropower, among all possible energy sector investments, are high. Confidence
56 that hydropower facilities will operate long into the future with performance at or near design
57 performance must be correspondingly high to justify investment. The 2016 Hydropower Status
58 Report of the International Hydropower Association (IHA, 2016) dedicates a chapter to the
59 subject and describes climate-specific resilience in three ways: 1) the ability to recover after an
60 external stressor or extreme event; 2) the capability to succeed in an environment dominated by
61 uncertainty; and 3) the capacity of a facility or system to withstand or adjust to the possible
62 impacts of climate change.

63 A number of studies have explored the climate change resilience of hydropower by
64 evaluating basin-wide changes in hydropower generation potential in the context of changes in
65 hydrology and water resources (e.g., Beyene, et al., 2010; Bharati, et al., 2014; Christensen, et
66 al., 2004; Christensen and Lettenmaier, 2007; Finger, et al., 2012; Giuliani, et al., 2016;
67 Grumbine, et al., 2012; Hamlet, et al., 2010; Ho, et al., 2016; Lehner, et al., 2005; Majone, et al.,
68 2016; Markoff and Cullen, 2008; Maurer, et al., 2009; Mehta, et al., 2011; Minville, et al., 2009;
69 Schaefli, et al., 2007), the ability of diminishing glaciers to continue to sustain baseflows on
70 which run-of-river hydropower facilities rely (Bolch, et al., 2012; Shrestha and Aryal, 2011), the
71 vulnerability of hydropower structures to glacier-lake outburst floods (Dussailant, et al., 2010),

72 and the impact of seasonality shifts on hydropower timing (Laghari, et al., 2012; Madani and
73 Lund, 2010; Sharma and Shakya, 2006). Some have found substantial evidence of the effects of
74 climate change on hydropower already: Destouni et al. (2013) in northern Europe; Hanshaw and
75 Bookhagen (2014) in the Andes, Peru; and Sorg et al. (2012) in Central Asia.

76 There are important limitations in the ability of these previous studies to inform risk-
77 management aspects of hydropower investment. All of the coupled hydrologic-hydropower
78 models cited in the previous paragraph used as climate input the output from general circulation
79 models (GCMs) from the Intergovernmental Panel on Climate Change (IPCC) Coupled Model
80 Intercomparison Project (CMIP), with the exception of Mehta et al. (2011), which used a
81 scenario of 2 degrees warming in the Sierra Nevada, California. Some assessed only streamflow
82 without investigating the facility itself (e.g., Ho, et al., 2016; Minville, et al., 2008), thereby not
83 identifying the vulnerabilities of hydropower to climate change in a systematic way. Even where
84 infrastructure models have been involved, the results have been contingent on the projections and
85 downscaling method that happened to be used. In many cases, these studies based their
86 conclusions regarding climate change vulnerability on model results forced with only one or two
87 climate change scenarios. By not systematically exploring climate change vulnerabilities, each
88 leaves unanswered the question of greatest concern to policy-makers grappling with the potential
89 risks and rewards of hydropower investment.

90 Furthermore, risks to hydropower investment are not limited to climate change. Recent
91 studies have found that capital cost overruns (Ansar, et al., 2014) and electricity selling price
92 (Gaudard, et al., 2016) are key concerns for hydropower investors. Other non-climate-change
93 risks are due to earthquakes, landslides, other natural disasters, or military action, with associated
94 risks of dambreak and flood surge to inhabitants and structures downstream (Benn, et al., 2012;

95 Dai, et al., 2005; Dussailant, et al., 2010; Peng and Zhang, 2012; Richardson and Reynolds,
96 2000), and storage loss from sediment accumulation (Annandale, 1987; Castillo, et al., 2015;
97 Wild, et al., 2016). It is clear that hydropower investment would benefit from a comprehensive
98 assessment of the uncertain factors that potentially impact the benefits and costs of hydropower
99 investments.

100 Previous studies have presented tools for multidimensional sensitivity analysis (Lempert,
101 et al., 2006; Lempert, et al., 2003) and applied those tools to water systems planning (e.g.,
102 Groves and Lempert, 2007; Kasprzyk, et al., 2013; Kwakkel, et al., 2016); however, those
103 studies are not targeted at hydropower, and none have demonstrated a multidimensional stress
104 test framework that addresses the shortcomings of GCM-led climate change risk assessments.
105 Groves et al. (2015) performed project-scale climate change vulnerability analysis on five
106 hydropower projects planned for sub-saharan Africa, and noted that the sensitivity of the
107 performance of two of the projects to hydropower selling price may be more significant than to
108 climate change, but did not evaluate the relative vulnerabilities quantitatively. Kucukali (2011)
109 presents a multidimensional risk assessment for hydropower projects that does not address
110 climate change risks, while Kubiszewski et al. (2013) presents a process for multidimensional
111 risk assessment of hydropower systems that gives only cursory attention to climate change
112 through the inclusion of a narrow set of prescribed climate change scenarios. Yang et al. (2016)
113 evaluated risks to the water-energy-food nexus of the Brahmaputra river basin, including a
114 thorough treatment of climate change risks, but did not address risks to hydropower investments,
115 in particular, or present a generalized methodology.

116 The process described in this paper assesses multidimensional risk to hydropower
117 investments including cost, selling price, discount rate uncertainty, natural hazards (e.g.,

118 landslide, earthquake), sediment damage to turbines, and a bottom up approach to climate change
119 risks. The process integrates simulated results from coupled climatic, glaciological and
120 hydrological models, informed by data from in situ and remote-sensing based measurements, that
121 are bottom-up and site-specific. To those elements is added an infrastructure model to evaluate
122 the resilience of hydropower facilities that is responsive to changes in both climate- and non-
123 climate factors. A stress-testing approach applied to the model chain, coupled with a data mining
124 algorithm, allows for identification of the relative significance of risks of different types to the
125 project. Once the project vulnerabilities are identified, adaptation options can be quantitatively
126 evaluated.

127 The manuscript is organized as follows: Section 2 describes the process, Section 3
128 demonstrates the process through an evaluation of a proposed hydropower project in the Arun
129 river basin of Nepal, and Section 4 presents areas for further research and concludes.

130 **2. METHODS**

131 The risk management framework presented in this paper was developed in response to a
132 recent mandate of the World Bank that all International Development Association (IDA) Country
133 Partnership Frameworks must include climate- and disaster-risk considerations in the analysis of
134 the country's development priorities, and, when agreed upon with the country, incorporated into
135 the content of the development programs. This mandate did not specify the means by which
136 climate change risks should be assessed, and no consensus existed on the appropriate process.
137 Context was provided, however, by the Independent Evaluation Group of the World Bank, which
138 found that "climate models have been more useful for setting context than for informing
139 investment and policy choices," and concluded that the prominent applications of climate change

140 projections to infrastructure performance analysis “often have relatively low value added for
141 many of the applications described” (IEG, 2012).

142 In response, a clear process for demonstrating the resilience of a water system investment
143 to climate, geophysical and economic uncertainty was presented in Ray and Brown (2015). The
144 process adopts bottom up decision scaling techniques (Brown, et al., 2012) for climate change
145 risk assessment, and provides guidance on methods for risk management. The process presented
146 in Ray and Brown (2015) is structured as a decision tree or decision flow that leads the analyst to
147 a particular analytical method based on the characteristics of the project being investigated. The
148 procedure consists of four successive phases: Phase 1 Project Screening; Phase 2 Initial
149 Analysis; Phase 3 Stress Test; and Phase 4 Risk Management. The project under investigation
150 moves through only as many phases of the process as are justified by need.

151 This study presents an implementation of this process for hydropower investments. All
152 hydropower projects are classified as “potentially climate sensitive” in Phase 1, and undergo
153 multifactor sensitivity analysis in Phase 2. If the multifactor sensitivity analysis shows climate
154 sensitivities to be significant relative to potential sensitivities of other kinds (e.g., financial,
155 natural hazard), then Phase 3 stress tests (climate and multidimensional) are initiated. The
156 concepts and methods of sensitivity analysis on which Phase 2 is built are described in Saltelli et
157 al. (2009). The insights gained in Phase 3 enhance the findings of Phase 2. As such, and in the
158 interest of brevity, only Phases 3 and 4 are discussed here.

159 It must be emphasized that every phase described in Ray and Brown (2015) is to be a
160 collaborative partnership with project investors and local stakeholders. Phases 1 and 2, in
161 particular, are consultative processes in which project objectives (performance criteria and
162 thresholds) are defined, and the set of risks against which the project performance is to be

163 evaluated is identified.

164 **2.1 Multidimensional Stress Test**

165 The multidimensional stress test places climate risks in the context of risks of other kinds.

166 The goal is to evaluate the response of the model workflow to a wide ranging set of uncertain

167 factors critical to system performance. The critical factors will generally include at least the three

168 non-climate risk factors identified in the introduction: cost uncertainty, risks to the structure of

169 the dam from natural disasters or military action, and sedimentation effects. These factors can be

170 efficiently sampled from ranges (or from probability density functions) defined by stakeholders

171 or through expert elicitation, using a Latin Hypercube algorithm (McKay, et al., 1979).

172 Exploratory analysis of efficiently sampled ranges of uncertain parameters has been the

173 backbone of a number of analytical techniques applied to decision support (e.g., Bankes, 1993;

174 Ben-Haim, 2006; Lempert, et al., 2006).

175 A climate stress test results when the multidimensional stress test is conducted while

176 holding all non-climate parameters constant. A climate stress test (Brown, et al., 2012) consists

177 of four steps: 1) the coordination of a workflow of models to translate climate inputs to outputs

178 descriptive of water system performance; 2) the use of the workflow of models to explore a wide

179 climate domain in order to empirically derive a “climate response function” or, if the function

180 happens to contain three dimensions, a “climate response surface”; 3) the development of

181 vulnerability-specific scenarios by parsing of the climate space according to problematic climate

182 conditions (i.e., the climate conditions in which system performance fails according to some pre-

183 defined threshold); 4) climate informed decision analysis in which likelihood concepts are

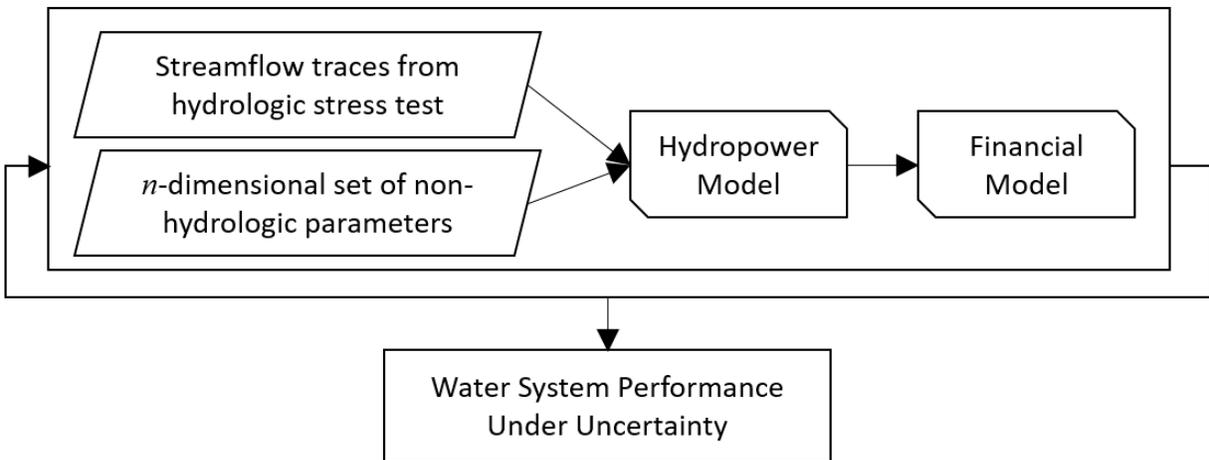
184 developed using the most credible information available, and climate change risks to the project

185 (or system) are evaluated.

186 The emphasis on vulnerability-specific scenario development and evaluation in the
187 climate stress test is in contrast to the more conventional “scenario-led” approach, in which the
188 performance of the system is evaluated using a sample of possible futures described by human
189 development storylines (taken, for example, from the IPCC special report on emissions scenarios
190 (Davidson and Metz, 2000) or its representative concentration pathways (RCP) update (Taylor,
191 et al., 2012)). In contrast to human development story-line-based approaches, climate change
192 scenarios can be generated by parametrically (Ben-Haim, 2006) or stochastically (Brown, et al.,
193 2012; Prudhomme, et al., 2010) varying the climate data to identify vulnerabilities in water
194 system performance, and elaborating scenarios according to the vulnerabilities of the project or
195 the opportunities it presents. This idea as applied to multidimensional stress tests is further
196 expounded upon in the Vulnerability-Specific Scenario Analysis section below.

197 In modeling hydropower performance, in particular, there are two broad categories of
198 vulnerability: hydrologic (encompassing climate-related parameters, land-use parameters, and
199 parameters related to the upstream consumptive water use or downstream water release
200 requirements), and non-hydrologic (financial, natural hazard, etc.). The workflow of hydrologic
201 stress test models includes: 1) a stochastic climate/weather generator to generate a wide range of
202 climate time series representative of historic natural variability and plausible climate shifts; and
203 2) a hydrologic model to translate climate traces into time series of available water. Permutations
204 of the streamflow traces generated during the execution of the hydrologic stress test can be input
205 directly to the hydropower model in combination with samples of the non-climate parameters on
206 which performance of the hydropower model depends. The multidimensional stress test (Figure
207 1) then inputs the hydrologic and non-hydrologic information into a hydropower simulation
208 model to translate time series of available water into estimates of hydro-electric power

209 production, and finally, a financial model that calculates financial performance across a range of
210 climate and socio-economic conditions. The arrow looping back from the output of the model
211 workflow to the input of the workflow implies iteration, in some cases thousands of times, to
212 trace out the project's vulnerability space.



213

214 Figure 1. Schematic representation of the multivariate stress test.

215 A brief description of each sub-step of the multidimensional stress test follows.

216 2.1.1 Stochastic Weather Generator and Hydrologic Model

217 Weather generators produce long series of synthetic weather data (Kilsby, et al., 2007;
218 Wilks and Wilby, 1999). As described in Steinschneider and Brown (2013), a weather generator
219 is conditioned on the historical data, and maintains the critical statistical properties of the
220 multivariate historical climate (e.g., low-frequency oscillations, temporal autocorrelations,
221 spatial correlations, mean, variance, skew). The weather generator then enables the systematic
222 perturbation of climate characteristics (mean shifts and changes in measures of variability) on
223 which the hydrologic stress test is built.

224 It is recommended that the weather generator: 1) be conditioned on historical climate, and
225 not the output of GCMs, in order to maintain the local characteristics of precipitation, in

226 particular, that GCMs do not faithfully reproduce (Rocheta, et al., 2014); 2) impose climate
227 changes (delta shifts, quantile mapping, or some more nuanced change in climate variability) that
228 far exceed that suggested by the current ensemble of GCM output. Justification for the second
229 recommendation includes the following reasoning: 1) the range of uncertainty in precipitation
230 change is larger in most cases in the CMIP5 ensemble than in the previous (CMIP3) ensemble;
231 2) all RCP scenarios from the IPCC's 5th Assessment Report assume a large reduction in the
232 atmospheric aerosol emissions by the end of the 21st Century, which is likely to be too narrow
233 (Stouffer, et al., 2017), and 3) any particular GCM ensemble would represent only the "lower
234 bound on maximum range of uncertainty" (Stainforth, et al., 2007).

235 The climate information from the weather generator is fed into a hydrologic model,
236 which translates climate variables to hydrologic variables of interest (e.g., streamflow at inflow
237 points to water infrastructure).

238 2.1.2 Hydropower Model

239 Streamflow traces output from the hydrologic model can be converted to kilowatt-hours of
240 electrical energy using:

$$241 \quad KWH_t = 0.002725 \cdot Q_t \cdot H_t / e \quad (1)$$

242 where Q_t = total streamflow through the turbines in time t [m³], H_t = net head in time t [m], and e
243 = the efficiency of the conversion of mechanical energy into electrical energy. The coefficient
244 0.002725 is an aggregate unit conversion. Detailed explanation of the derivation and utility of (1)
245 is available in Loucks and van Beek (2005).

246 In the case of run-of-the-river hydropower, H is typically constant. In storage reservoir
247 applications, H is a function of the reservoir storage level.

248 2.1.3 Financial Model

249 With the effect of climate change on local streamflow established by way of a hydrologic
250 model, and the effect on hydro-electric production established by way of a hydropower model,
251 the inclusion of a financial model in the multidimensional stress test allows the examination of
252 the effect of change on the financial performance of the hydropower system. Whenever possible,
253 it is important to consider distribution equity, and as wide a range of social and environmental
254 costs and benefits as possible (Ansar, et al., 2014; Evans, et al., 2009; Rosenberg, et al., 1995),
255 though time and data limitations do not always allow for the preferred depth of economic
256 analysis of hydropower projects. In such cases the economic model may be limited to strictly
257 financial considerations.

258 There are a number of representations of the financial value of hydropower investments
259 available, such as life cycle cost (e.g., Zakeri and Syri, 2015), levelized cost (e.g., Jaramillo, et
260 al., 2004), and internal rate of return or net present value (e.g., Mishra, et al., 2011; Santolin, et
261 al., 2011). Financial models, as a subset of economic models, are most applicable when
262 evaluating the system from the perspective of the owner or investor, as is the case here.

263 **2.2 Climate Informed Risk Assessment**

264 In this step, the full ensemble of available climate information (e.g., local historical
265 trends, paleo records, global climate change projections) is used to inform the likelihood of the
266 types of climate changes to which the project is shown vulnerable in the multidimensional stress
267 test. To do so, the following steps are involved: 1) a multi-model, multi-realization ensemble of
268 all available (or all relevant) GCM projections is assembled; 2) GCM projections are assessed in
269 terms of their biases and ability to credibly reproduce the relevant climate variables (Bush, et al.,
270 2015; Wang, et al., 2014; Wilby and Harris, 2006), which may result in the need for corrective

271 downscaling (Wilby and Dawson, 2013); and 3) projections are superimposed on the climate
272 response surface, providing an indication of the likelihood of vulnerability-specific scenarios
273 developed in step 3 of the climate stress test (Manning, et al., 2009; Steinschneider, et al., 2015).

274 Likelihood functions developed using GCM output can range from simple GCM-count
275 approximations (e.g., Brown, et al., 2012) to Bayesian approaches to determining probability
276 distribution functions of climate change at regional scales (e.g., Tebaldi, et al., 2005) to more
277 complex hierarchical Bayesian models that generate summary bivariate Gaussian probability
278 density functions that account for correlations between non-independent GCM output (e.g.,
279 Steinschneider, et al., 2015).

280 **2.3 Vulnerability-Specific Scenario Analysis**

281 The output of the multidimensional stress test is likely to be large, complex, many-
282 dimensional, and difficult to visualize. Interactive versions of many-dimensional scatterplots
283 presented in three-dimensional space or interactive parallel coordinate plots are useful
284 approaches to many-dimensional data visualization. See Kasprzyk et al. (2013) and Yang et al.
285 (2016) for examples.

286 Beyond multidimensional graphical visualization of the output, computational data
287 mining tools are typically required to identify problematic combinations of change in the set of
288 uncertain future conditions. Any cluster analysis technique (e.g., density-based or shared-
289 property-based) might be appropriate (see Tan et al. (2005) for examples). The case study that
290 follows uses the Patient Rule Induction Method (PRIM) (Friedman and Fisher, 1999), a
291 statistical cluster-finding algorithm that identifies combinations of conditions that result in
292 system failure relative to a pre-defined performance threshold.

293 **3. DEMONSTRATION OF APPROACH**

294 **3.1 Background on the Upper Arun Hydropower Project**

295 A land-locked, low-income country endowed with rich hydropower resources, Nepal
296 views hydropower development as its key opportunity for economic growth and human
297 development (Bonzanigo, et al., 2015). It is estimated that Nepal has approximately 40,000 MW
298 of economically feasible hydropower potential, of which only approximately 600 MW has so far
299 been developed (World Energy Council, 2013).

300 Because of its large reliance on hydropower (hydro-electricity generation comprises 90%
301 Nepal's current national energy generation portfolio (IDS-Nepal, 2014)), the energy-generating
302 capacity of Nepal is particularly low in the dry season when monsoon flows and glacier melt are
303 not abundant. Demand is fairly steady throughout the year, resulting in a mismatch in the
304 seasonality of energy supply and demand in Nepal (NEA, 2014).

305 Much of the hydropower development planned for Nepal is concentrated in the Koshi
306 River basin (Chinnasamy, et al., 2015; Hosterman, et al., 2012). The Upper Arun Hydropower
307 Project (UAHP) on the Arun River in the Koshi basin represents one of Nepal's highest priority
308 hydropower development projects. The Arun River originates at a glacier on the northern slope
309 of Mount Xixabanma in Tibet. The river flows approximately 300 km eastward across the
310 Tibetan Plateau at an elevation of about 4000 m before crossing the Himalayas and plunging into
311 Nepal where it flows in a narrow 30 to 60 m canyon at a very steep slope. The proposed dam site
312 is in the steep narrow, approximately 15 km downstream of the border with Tibet. The river has
313 the catchment area of nearly 25,700 km² (only 400 km² of which is within Nepal) and an average
314 run-off approximately 200 m³/s. Very little development has so far occurred in the catchment
315 area (Latham, et al., 2014).

316 The original feasibility study (NEA, 1991) resulted in the recommendation of an
 317 installed capacity of 335 MW in a peaking run-of-river (PROR) scheme. The design discharge
 318 was 78.8 cubic meters per second (m³/s), Q70 (streamflow sufficient to operate turbines as
 319 design capacity 70% of the time), and it was expected to generate 2050 gigawatt-hours (GWh)
 320 per year. The estimated capital costs when the Nepal Electricity Authority (NEA) obtained
 321 license for the project in 2008 was US\$446 million. UAHP baseline specifications are
 322 summarized in Table 1. Values in Table 1 were obtained during a stakeholder workshop hosted
 323 in September 2013 in Kathmandu by the World Bank, and joined by project stakeholders
 324 including local NGOs, academics, and the hydrologic, climate and energy services of the
 325 Government of Nepal (GoN).

326 Other, larger design sizes (750 MW, Q40, US\$1.01B; 1000 MW, Q35, US\$1.35B; 1355
 327 MW, Q30, US\$1.82B; 2000 MW, Q25, US\$2.69B) are also being considered as adaptation
 328 options for the UAHP, and will be discussed in reference to the multidimensional stress test later
 329 in this manuscript.

330 Table 1. UAHP baseline specifications

Parameter	Value
Capital cost	US\$450M
O&M cost	$250000 \cdot (installed_cap)^{0.65}$
Energy gen. cap.	335 MW (8.04 GWhr/day)
Plant load factor	0.75
Discount rate	5%
Dry season selling price	US\$0.084/kWh
Wet season selling price	US\$0.045/kWh
Economic lifetime	30 years (5 yrs construction)

331
 332 The O&M cost relationship is derived empirically from data on operation and
 333 maintenance costs in previous installations in Nepal (and other similar environments worldwide),
 334 and the parameter *installed_cap* is the installed capacity of the hydropower facility in units of

335 megawatts (MW). Plant load factor is here taken to be the plant availability considering shut
 336 down for sediment management. For example, a plant load factor of 0.75 means that the plant is
 337 shut down one quarter of every month for management of sediment (e.g., flushing) and other
 338 purposes.

339 Table 2 presents the baseline value for each uncertain factor and the expanded range of
 340 analyzed values. Ranges are informed by evaluation of locally-relevant data, and by consultation
 341 with stakeholder experts, and do not bound the *possible* universe of values. Instead, selected
 342 ranges are intended to be wide enough that no *plausible* risks are missed. The timeframe of the
 343 analysis includes only the project’s economic lifetime, extending approximately 30-40 years into
 344 the future.

345 Table 2. Uncertainty ranges

Uncertainties	Baseline value	Range min	Range max
1. Natural System			
Precipitation, change from historic mean	0%	-40%	+40%
Temperature, change from historic mean (°C)	0	0	+8
Plant load factor (surrogate for sediment effect)	0.75	0.60	0.90
Project lifetime (surrogate for seismic/landslide risk)	30	15	36
2. Electricity Markets and Prices			
Wet season electricity price, Apr-Oct (US\$/kWh)	0.045	0.045	0.135 (baseline x3)
Dry season electricity price, Nov-Mar (US\$/kWh)	0.084	0.084	0.252 (baseline x3)
3. Project Variables			
Capital costs (year 2013 US\$M)	446	446	1338 (baseline x3)
Discount rate	0.05	0.03	0.12

346
 347 Delays in the initiation of construction of hydropower projects can lead to extreme cost
 348 overruns. Construction delays resulted in a tripling of the final capital costs for the Marshyangdi
 349 Dam in Nepal, for example. A correspondingly large consequence of construction delays was

350 therefore applied to the UAHP; this was represented in the model by sampling from a range of
351 capital costs from the baseline to up to three times the baseline. Electricity prices are subject to a
352 similar magnitude of uncertainty. If the GoN were to begin exporting electricity to India, for
353 example, the selling price would be on the order of 0.15 US\$/kWh, or approximately triple the
354 current wet season selling price. The baseline plant load factor is understood to be approximately
355 0.75, meaning that the turbines are shut down for cleaning and flushing 25% of the time in
356 normal operation. Higher sediment rates than expected could require turbine off-times to
357 increase to 40%; lower sediment rates could drop turbine off-times to 10%. Earthquakes,
358 landslides, and military action could cause damage to the structure of the hydropower facility
359 and shorten its useful lifetime. As a surrogate for damage to the structure from natural (or
360 military or other) hazards, the possibility that the 30-year baseline lifetime could be cut short (by
361 up to half) was modeled; and the model was also tested with longer-than-expected economic
362 lifetimes up to an additional 6 years of operation was also considered. If properly managed,
363 hydropower investments could last much longer than 30-40 years; however, when even a low-
364 value discount rate is applied, the NPV is not appreciably affected by extensions of design life
365 beyond 30 years.

366 Assignment of the discount rate is a political process, often contested (e.g., Mendelsohn,
367 2008; Nordhaus, 2007), and influences how resources are allocated between the present and the
368 future (Arrow, et al., 2004). A higher discount rate signifies an urgency to satisfy present
369 needs; whereas a lower discount rate values consequences of the present investment (positive
370 and negative) further into the future. In practice, the social discount rates used to evaluate the net
371 benefits of proposed projects have varied widely, with developed nations typically applying a
372 lower rate (3–7 percent) than developing nations (8–15 percent) (Zhuang, et al., 2007).

373 Organizations such as the World Bank, the Inter-American Development Bank, and the Asian
374 Development Bank use a discount rate of 12 percent, or in some cases (such as water supply
375 projects), 10 percent. Recently, however, discussions have turned to the consideration of lower
376 discount rates (3-6%) that better represent social welfare (World Bank, 2016). The essential
377 rationale for these elevated discount rates stems from the high value of scarce capital in
378 developing countries: projects consuming large amounts of capital are required to account for the
379 opportunity cost of these financial resources, pushing up the expected rate of return (Goulder and
380 Williams III, 2012). A range of discount rates from 3% to 12% was used in the UAHP analysis
381 to explore the effect of discounting on considerations of the sustainability of the investment.

382 Climate change is a prominent concern to all involved in the UAHP planning process.
383 Climate change everywhere is difficult to anticipate, but in the region of the Himalayas,
384 prediction of future climate is particularly difficult (Nepal and Shrestha, 2015; Singh and
385 Bengtsson, 2004). There is great uncertainty in climate projections due to the complex
386 topography, the importance of the South Asian Monsoon (Loo, et al., 2015; Turner and
387 Annamalai, 2012), and uncertainty associated with glacier volumes (Savoskul and Smakhtin,
388 2013). Historical measurements are sparse due to the high elevations and forbidding terrain. The
389 historical analysis that has been accomplished shows increasing temperatures and no clear signal
390 on precipitation (Akhtar, et al., 2008; Immerzeel, et al., 2013; Immerzeel, et al., 2010). Glaciers
391 are growing in some areas and receding in other areas (Bajracharya and Shrestha, 2011; Kaab, et
392 al., 2012), and more are receding than growing. Streamflow generally seems to be increasing
393 (Bookhagen and Burbank, 2010; Lutz, et al., 2014).

394 **3.2 Model Workflow**

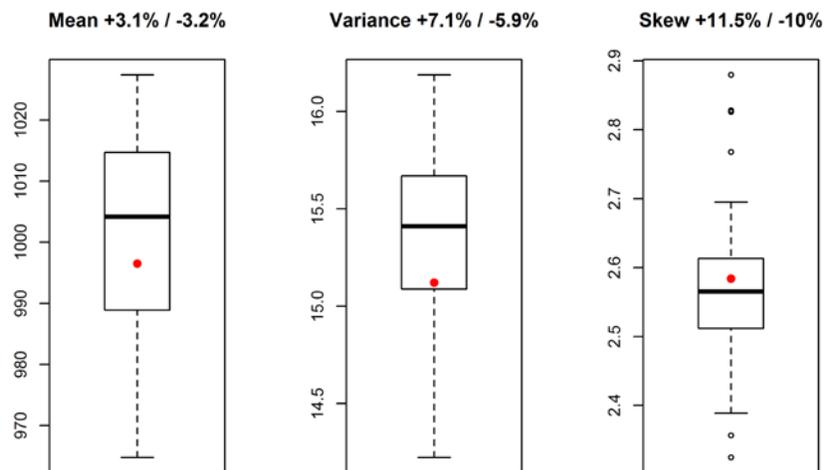
395 3.2.1 Stochastic Weather Generator

396 The climate change aspects of the multidimensional stress test presented in this section
397 has aimed to account for most, though not all, of these climate-related uncertainties, focusing on
398 those for which the credibility of projections is higher. Limitations related to climate-change
399 induced exacerbation of precipitation extremes are presented in the discussion section.

400 The weather generator resampled from 60 (1950-2010) years of 0.5 degree latitude and
401 longitude gridded daily temperature and precipitation data from the APHRODITE database
402 (Yatagai, et al., 2012). APHRODITE precipitation data were bias-corrected using GPCP data
403 (Schneider, et al., 2014). No adjustments were made to APHRODITE temperature data.

404 A wavelet transform analysis conducted for the basin-wide average annual precipitation
405 found no statistically significant (at the 90% confidence level) low frequency oscillation in the
406 historical record. Without need to preserve semi-oscillatory behavior of the long-term
407 precipitation signal, and given both 1) low confidence in the historical climate observations in
408 the Himalayan region (National Research Council, 2012), and 2) the interest of the project
409 stakeholders in simple, direct techniques for the generation of climate traces, a weather generator
410 simpler than others that have been used for a similar purpose (Groves, et al., 2008;
411 Steinschneider and Brown, 2013; Yates, et al., 2003) was employed in this case. The algorithm
412 developed for this study bootstrapped on the historical climate record (e.g., Vogel and
413 Shallcross, 1996), for all APHRODITE grid cells of the basin simultaneously (to maintain spatial
414 correlations), by monsoon/non-monsoon season (to maintain temporal correlations). A sample of
415 thirty 36-year (the upper end of the range of anticipated economic lifetime of the hydropower
416 investment, discussed further in reference to the multidimensional stress test below) historically-
417 representative climate traces were developed. Annual precipitation statistics for the bootstrapped

418 traces are shown in Figure 2. Delta shifts were then applied to precipitation and temperature in
419 order to explore the effects of climate change (0-8 degrees Celsius increase in temperature and
420 $\pm 40\%$ change in precipitation relative to historic values, each applied uniformly throughout the
421 year). The range was selected in order to evaluate a range of climate changes well beyond what
422 is projected to occur in the watershed according to the full ensemble of CMIP5 GCM
423 projections. The range is intended to be wide enough that no precipitation- or temperature-shift-
424 related vulnerabilities are overlooked.



425
426 Figure 2. Annual precipitation statistics for the thirty 36-yr climate traces ($n=30$) of the
427 UAHP weather generator: mean, variance, and skew. The red dot indicates the statistics of
428 the historical record (1950-2010).

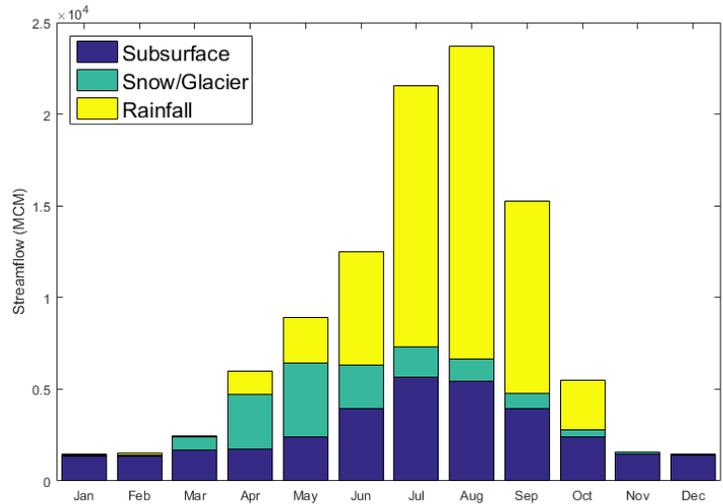
429 3.2.2 Hydrologic Model

430 In this study, we adopt the distributed glacio-hydrologic model HYMOD_DS (Wi, et al.,
431 2015), which was developed specifically for mountainous regions with sparse data. The model is
432 composed of hydrological process modules that represent soil moisture accounting,
433 evapotranspiration, snow processes, glacier processes and flow routing. The model operates on a
434 daily time step and requires daily precipitation and mean temperature as input variables. The

435 overall structure of the HYMOD_DS is described in Wi et al. (2015). The snow/glacier module
436 is critical for the test basin in this study where a snow/glacier melting water is dominated source
437 of water. Depending upon the size of the basin, HYMOD_DS calibration can be time-intensive,
438 and may be expedited with the aid of parallel computing power.

439 The streamflow data used for calibration were obtained from the Nepal Department of
440 Hydrology and Meteorology, site 600.1, Uwagaon, with coverage from April 1986 through
441 December 2006. A Nash Sutcliffe Efficiency value of 0.91 was achieved for the calibration
442 phase (1986-1995), and with a corresponding value of 0.75 for the validation phase (1996-2006).

443 Figure 3 presents output of the hydrologic model showing the component contributions of
444 subsurface groundwater flow, snow/glacier melt, and rainfall runoff to streamflow at Uwagaon
445 station. The historical time series shows little change in glacier/snow contribution. The seasonal
446 hydrograph shows that the greatest contribution of meltwater occurs in April/May/June, and the
447 greatest contribution of rainfall runoff occurs July/August/September. Streamflow from
448 November through February is supported almost exclusively by groundwater baseflow.
449 However, it must be noted that no ground-truth measurements of the relative contributions to
450 streamflow of groundwater or snow/glacier, or even reliable measurements of changes in aquifer
451 storage or snow/glacier depth, were available for calibration of the glacio-hydrologic model. The
452 values presented in Figure 3 are therefore based on surface-groundwater equilibrium equations
453 and timeseries of change in glacier area (Randolph Glacier Inventory version 3.2, RGI 3.2,
454 Pfeffer et al. (2014)).



455

456 Figure 3. Hydrologic model calibration results showing percent contribution to streamflow

457 Figure 4 illustrates the long-term average annual, dry season, and wet season streamflow

458 at Uwagaon station, subjected to a range of climate conditions. Precipitation has the dominant

459 effect on streamflow, as demonstrated by the largely vertical contour lines. Streamflow shows a

460 more or less monotonic response to changes in precipitation, i.e., increases in precipitation result

461 in increases in streamflow and decreases in precipitation result in decreases in streamflow.

462 Temperature effects are smaller but more interesting. Over the course of the 30-year simulation,

463 a critical inflection point in the flow pattern occurred at an increase in temperature of

464 approximately 3 degrees C (less in runs with strong negative precipitation shifts). When the

465 system was simulated with temperature increases less than 3 degrees C over the historic, the

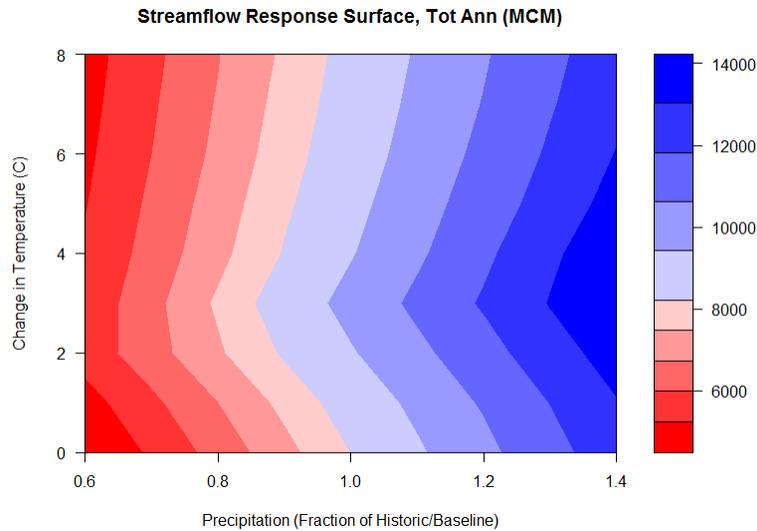
466 increased temperature exhibited a positive effect on streamflow resulting from greater quantities

467 of meltwater contribution from snow/glacier. However, with temperature increases larger than 3

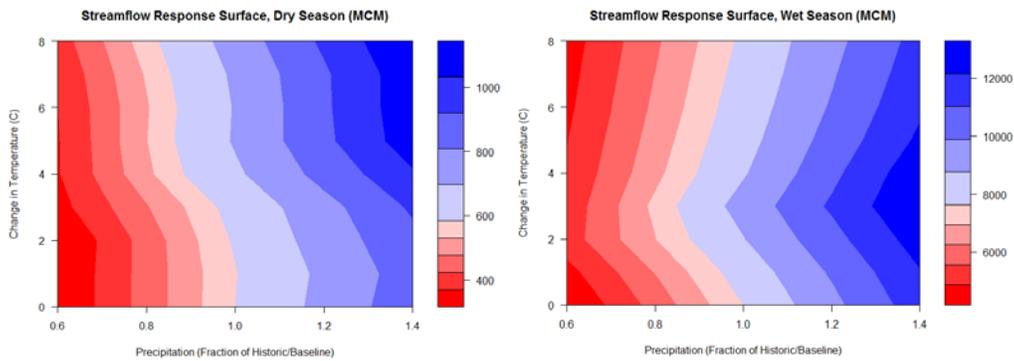
468 degree C, the streamflow gains are reversed as increasing rates of evapotranspiration and

469 diminishing returns from a shrinking (receding) glacier decreased the total rate of flow. This

470 phenomenon is especially evident in the wet-season response, as most of the meltwater is
471 contributed after March, the final month of the dry season (November-March).



472



473

474 Figure 4. Response of streamflow to changes in climate by annual total, dry season (lower
475 left) and wet season (lower right). Changes in precipitation are shown on the x axis and
476 changes in temperature are shown on the y axis. Contour colors represent increasing
477 streamflow, blue in excess of historic mean, and red less than historic mean.

478 The differences between the dry season and wet season in Fig 5 are partly attributable to
479 changes in precipitation type, and partly a result of glacier meltdown. As temperature increases,
480 winter (dry season) flow increases (as precipitation that falls in the winter is not stored as snow,

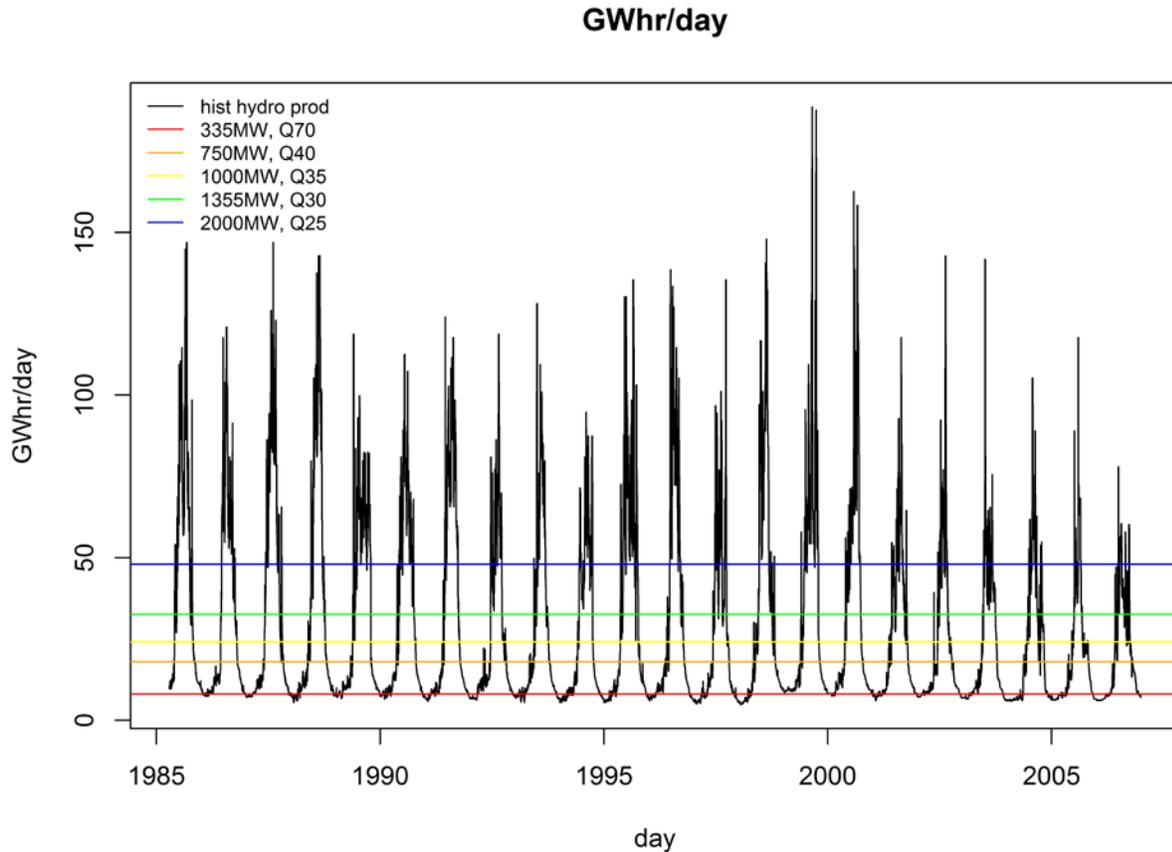
481 but runs off immediately into the stream). Peak flow also increases due to enhanced glacial melt.
482 There is a shift in peak meltwater contribution from Apr-May to Mar-Apr. The analysis is based
483 on glacier coverage map data obtained from RGI 3.2 (Pfeffer, et al., 2014), and glacier volume
484 was estimated using the multivariate glacier area-volume scaling relationships proposed by
485 Grinsted (2013). Temperature is the dominant factor in the recession of the glacier area. Because
486 a 3 degree C temperature increase throughout the 30-year simulation reduces the glacier volume
487 to 60-70% (relative to the initial volume), the remaining glacier area/volume is insufficient to
488 continue to sustain streamflow at historic levels. The significance of this result is that more
489 streamflow may be available when it is most needed to produce high-value dry-season
490 hydropower.

491 3.2.3 Hydropower Model

492 Figure 5 shows streamflow data from Nepal Department of Hydrology and Meteorology,
493 site 600.1, Uwagaon, converted to GWhr/day using (1), with Q in units of million cubic meters
494 per day [MCM/day], $H = 492$ m (constant H , as the UAHP is a run-of-river project), and $e = 0.9$.
495 The horizontal red line in Figure 5 locates the capacity of the baseline 335 MW facility.
496 Hydropower production potential in excess of the red line would not be generated with a 335
497 MW facility, implying there would be much potential untapped.

498 The hydropower model was written in the R modeling environment. The performance of
499 the suite of increasingly large design adaptation alternatives (750 MW, 1000 MW, 1355 MW,
500 and 2000 MW) was analyzed using the same methodology in order to understand the extent to
501 which those designs could better capitalize on increased streamflow from increased glacier melt
502 and potentially greater monsoon rainfall, as well as to understand the financial risks posed by
503 investment in larger hydropower facilities. Hydropower production potential for these other

504 proposed UAHP design sizes are also shown on Figure 5, and will be discussed in reference to
 505 risk management in a later section.



506
 507 Figure 5. UAHP daily GWhr/day generation relative to daily time series of GWhr potential.
 508 Horizontal lines represent limit of daily generating for each considered UAHP design size.

509 3.2.4 Financial model

510 Time and data limitations did not allow for the development of the broader economic
 511 considerations of this project, and so a strictly financial model was used to represent the costs
 512 and benefits of the project to the investor. The NPV was calculated using (2):

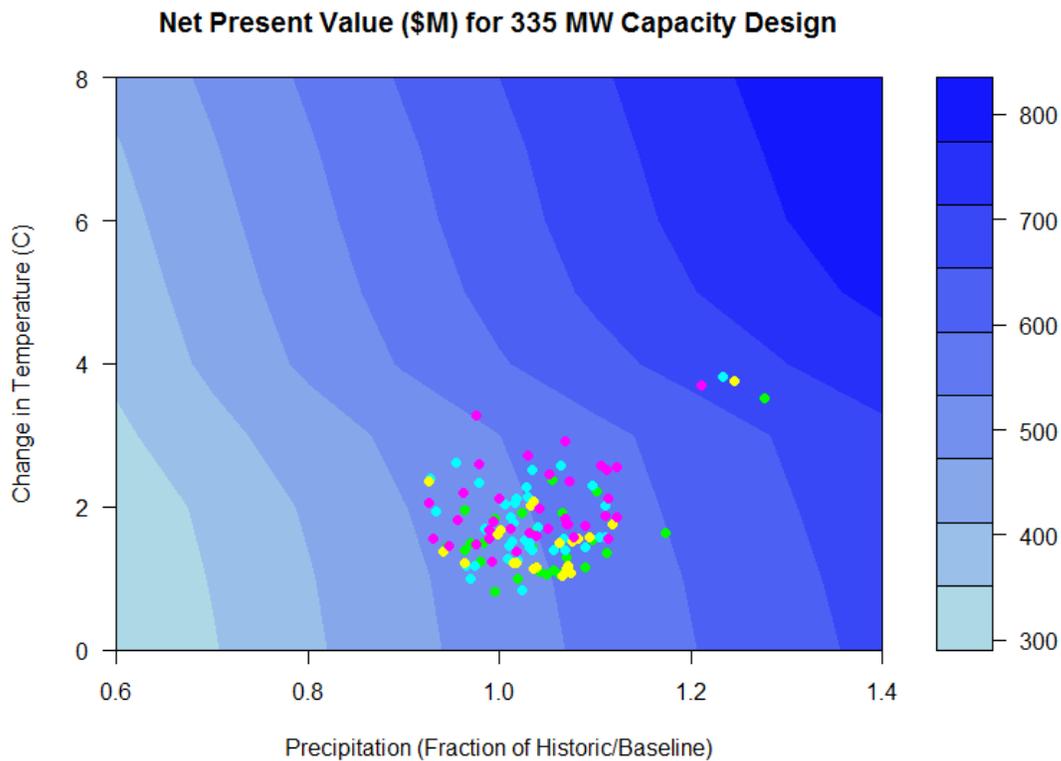
513
$$NPV = \sum_{i=1}^T \frac{C_t}{(1+r)^t} - C_0 \quad (2)$$

514 where C_t = net cash inflow during the period (hydro-electricity sales minus operating costs
515 [US\$M]); C_0 = initial investment [US\$M]; r = discount rate; and t = number of time periods
516 [months].

517 **3.2 Results**

518 3.2.1 Climate Informed Risk Assessment

519 Looking first at the impact of climate changes on the performance of the project, a
520 response surface presenting the climate change effect on the NPV of the baseline (335 MW)
521 hydropower investment is shown in Figure 6. Figure 6 was developed with Table 1 baseline
522 values for capital cost, O&M cost, turbine capacity, plant load factor, discount rate, selling price
523 (dry season and wet season), and economic lifetime. Using Table 1 baseline values for all non-
524 climate parameters, the NPV is positive over the entire range of the climate stress test with
525 values of US\$300M-US\$800M. Hydropower system performance for non-climate parameter
526 values other than the baseline is explored in reference to risk management later on.



527

528 Figure 6. NPV of the 335 MW design with CMIP5 climate change projections (centered on
 529 year 2050) superimposed. Green: RCP 2.6; Cyan: RCP 4.5; Yellow: RCP 6.0; Magenta:
 530 RCP 8.5

531 Figure 6 shows that for the wide range of climate changes considered, the baseline UAHP
 532 design has a positive NPV. This suggests that the project is robust to climate changes. For
 533 additional information related to the plausibility of the specific climate changes explored, mean
 534 changes from the full ensemble of CMIP5 GCMs are provided.

535 Climate projection-based analyses typically attempt to infer climate changes from model
 536 projections by ignoring biases and calculating the change between the historical simulation and a
 537 future projection that incorporates increasing greenhouse gas emissions. However, if the changes
 538 that are calculated are smaller than the biases, it is difficult to infer a direction of change since
 539 the changes are within the range of the model errors (and thus cannot be separated from noise).

540 In the case of precipitation output of the full ensemble of CMIP5 GCMs overlaying the UAHP
541 watershed, the average bias (1971-2000) is approximately 50% and the average change (2036-
542 2065 relative to 1971-2000) is approximately 1%. Given that note of caution, the changes
543 derived from the projections are calculated. On average, the multi-model, multi-run ensemble of
544 GCMs projections show no clear signal in terms of precipitation change. Temperature
545 projections generally show 1-4 C increases in temperature, as would be expected. More warming
546 is projected for the winter than the summer.

547 Figure 6 brings together the results of the climate stress test with the mean climate
548 changes derived from raw CMIP5 GCM climate projections for the basin. The figure shows that
549 for most climate projections, the NPV is little changed from the baseline estimate, and in many
550 cases, has higher NPV. There are few cases where the NPV is lower than the baseline estimate.

551 In general, the results can be interpreted as not providing any strong concerns that there
552 are problematic climate changes expected, given the response to changes in climate shown in
553 Figure 4, and most important, the consistently positive NPV shown in Figure 6 for the widest
554 range of climate changes. The results presented here are averaged across 30 realizations of
555 internal climate variability described in the description of the weather generator above.

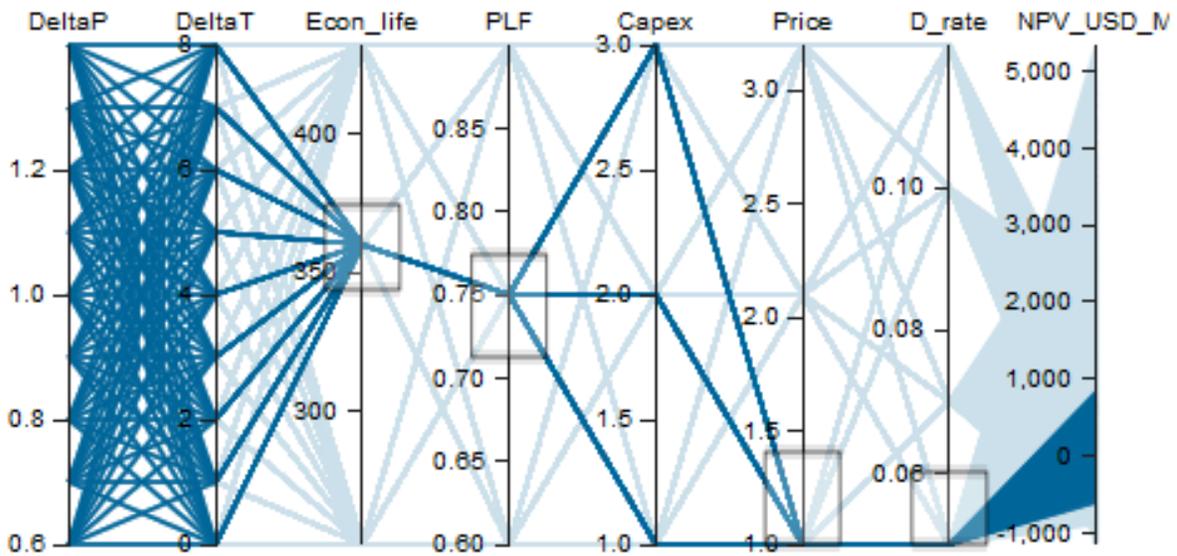
556 In summary, the 335 MW pre-feasibility design appears to be robust to climate changes.
557 However, the effect of climate change in combination with other uncertainties may reveal
558 problematic scenarios. For a full assessment of the robustness of the project design, these non-
559 climate factors are added to the analysis.

560 3.2.2 Vulnerability-Specific Scenario Analysis

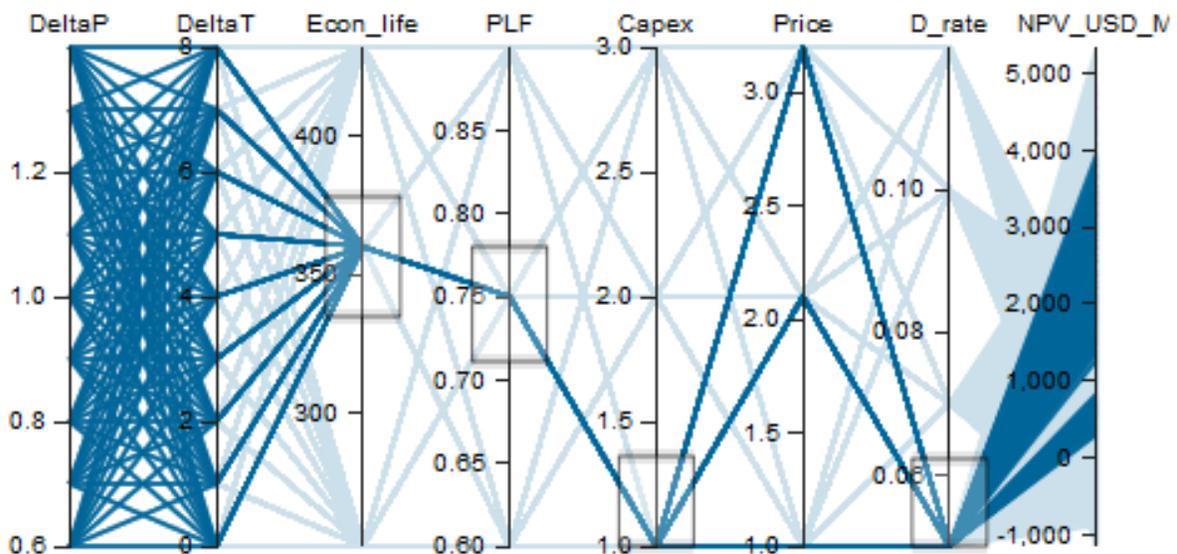
561 The sample of thirty historically-representative climate traces produced by the weather
562 generator were averaged, and 81 climate change factors (9 temperature shifts x 9 precipitation

563 shifts) were applied as described in the weather generator section above. In the interest of
564 computational efficiency, of those 81 traces, 13 were selected in order to cover the climate
565 change space, and span the range of mean and variance of streamflow. The total number of
566 scenarios developed to test the performance of the 335 MW design was then 6500: 500 socio-
567 economic scenarios (combinations of non-climate uncertainties sampled using a Latin Hypercube
568 sampling technique) combined with the 13 “representative” hydrologic sequences.

569 The NPV of the 335 MW UAHP project is positive in the vast majority of the generated
570 futures. As shown in Figure 6, the baseline NPV under historical climate conditions is
571 approximately US\$500M, and the uncertainty in climate change extends the NPV results
572 throughout the range from approximately US\$300M to US\$800M. Figure 7 presents a parallel
573 coordinates plot that expands Figure 6 to include other dimensions of uncertainty, in addition to
574 climate change uncertainty. The column names of Figure 7 are: DeltaP = change in precipitation
575 relative to the historical (%); DeltaT = change in temperature relative to the historical (oC);
576 Econ_life = expected economic lifetime of the project (months); PLF = plant load factor, the
577 fraction of time the turbines operate when not being flushed for sediment buildup; Capex = the
578 capital cost of the project as a fraction of the baseline cost; Price = the selling price of electricity
579 as a fraction of the baseline price; D_rate = the discount rate; NPV_USD_M = the net present
580 value of the project in millions of US dollars. Figure 7(a) shows that capital cost uncertainty,
581 when added to climate change uncertainty, expands the range of uncertainty in NPV downward
582 to -US\$600M, with 40% of all possible scenarios falling below zero NPV. Figure 7(b) shows,
583 alternatively, that price uncertainty expands the climate change uncertainty space upward to
584 US\$4 billion, with 67% of all possible scenarios now rising above an NPV of US\$1 billion.



585



586

587 Figure 7. Parallel coordinates plots summarizing effect on NPV of uncertainty in capital
 588 costs (top – NPV range -US\$600M to US\$800M) and electricity price (bottom – NPV
 589 range US\$300M-US\$4B).

590 The data mining algorithm selected for this analysis, PRIM, focuses the attention of
 591 project planners on the uncertain parameters most relevant to questions of future project
 592 performance. This step is an integral part of the Robust Decision Making (Lempert, et al., 2006)

593 methodology, which it uses to define policy-relevant scenarios (Groves and Lempert, 2007)
594 useful to policy-makers debating the risks faced by a set of project investment alternatives.

595 PRIM uses four measures to evaluate the different sets of conditions it identifies:
596 coverage (the proportion of all failure conditions captured by the vulnerability-specific scenario),
597 density (the proportion of all conditions captured by the vulnerability-specific scenario that are
598 failures), support (the “size” of the vulnerability-specific scenario relative to the entire
599 uncertainty sample), and interpretability (the number of uncertain conditions used to define the
600 vulnerability-specific scenario; a measure of the ease with which the scenario can be understood
601 and communicated to policy-makers and stakeholders). PRIM allows the generation of tradeoff
602 curves that help users choose explanatory scenarios with the best combination of density,
603 coverage, support and interpretability.

604 Through Figure 7 we are able to identify the relative magnitude of failure zones resulting
605 from each evaluated uncertainty set. Parallel coordinates plots like that shown in Figure 7 can be
606 used interactively to isolate the particular combinations of inputs that *always* result in water
607 system performance above or below a threshold. However, such plots do not provide nuanced
608 information about the types of conditions that lead to performance successes or failures.
609 Vulnerability-specific scenario analysis informs not only system sensitivity to uncertainty in
610 input parameters, but also the co-occurrence of conditions that together describe a scenario in
611 which the project *typically* (though not always) carries a negative NPV, as summarized in Table
612 3, with results now presented for all versions of the UAHP design.

613 Table 3. Selected vulnerability-specific scenarios for each UAHP design alternative

Uncertain Parameter	335 MW	750 MW	1000 MW	1355 MW	2000 MW
Mean annual precipitation	< 0.7 x historical	--	< 1.1 x historical	--	--
Mean temperature increase	--	> 2.5°C	--	--	--
Lifetime of investment	--	--	--	< 1.1 x baseline	< 1.133 x baseline
Plant Load Factor	--	--	--	--	--
Capital cost	> 2x baseline	> 1.9 x baseline	> 1.8x baseline	> 1.5 x baseline	> 1.5 x baseline
Wet season electricity price	< US\$0.08 / kWhr	< US\$0.8 / kWhr	< US\$0.10 / kWhr	< US\$0.10 / kWhr	< US\$0.11 / kWhr
Discount Rate	--	--	--	--	--
<i>Coverage</i>	29%	57%	54%	72%	70%
<i>Density</i>	82%	81%	80%	80%	88%

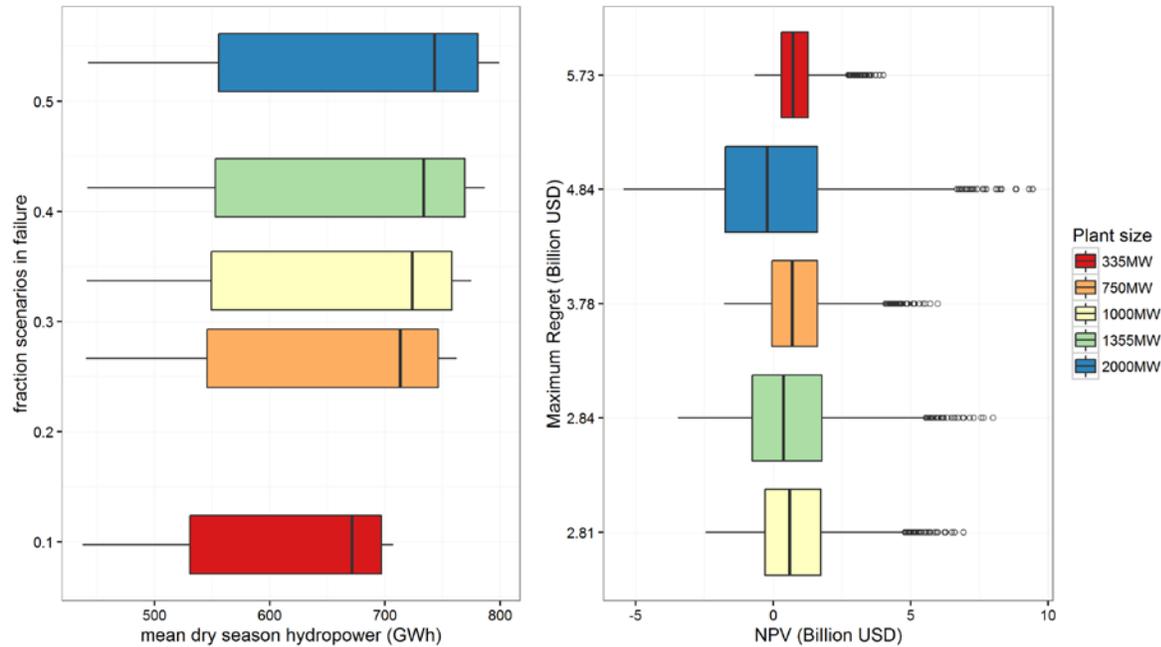
614

615 The multi-dimensional stress test of the 335 MW UAHP design capacity indicated that,
616 of all relevant uncertainties, the project is most vulnerable to high capital costs (more than twice
617 the baseline), low electricity prices (less than the wet season baseline), and decreasing
618 precipitation (more than 30% less than historical). When viewed in light of the low likelihood of
619 the co-occurrence of the three conditions of concern, however, the vulnerability of the project
620 appears small. There is only a small risk of such a marked decrease in precipitation (Figure 6
621 shows that no part of the ensemble of CMIP5 GCM projections indicates that the mean annual
622 precipitation would decrease by more than 30%) co-occurring with both a doubling of capital
623 cost and a failure of the selling price of electricity to rise above US\$0.08 (as would likely occur
624 were India and Nepal to agree to the electricity-sharing treaty).

625 For each of the design alternatives, NPV is more sensitive to capital cost and wet season
626 electricity selling price than to almost any other factor (climate change, included). Larger capital

627 investments (1355 MW and 2000 MW alternatives) are more sensitive to design life, requiring a
628 longer timeseries of steady hydropower sales to justify the construction costs.

629 Figure 8 presents tradeoffs in the performance of the five evaluated designs, and climate
630 change sensitivities provide an important insight regarding the timing of hydropower production.
631 The climate change stress test undergirding this multidimensional stress test accounts for the
632 climate-change-induced shift in peak meltwater contribution from April-May to March-April. In
633 the wet season, the 2000 MW design could produce up to 5 times more electricity than the 335
634 MW design; however, in the dry season (Figure 8(a)), the two are similarly limited by low flow,
635 with median dry season hydropower production for the 2000 MW design exceeding that of the
636 335 MW design by only approximately 10%. Furthermore, Figure 8(a) shows that the 10%
637 increase in median dry season hydroelectricity production is won at a cost: the 2000 MW facility
638 fails to meet its financial objectives in over 50% of the model runs, compared to only 10% for
639 the 335 MW design. Because wet season hydropower is of low value relative to dry season
640 hydropower, this comparative advantage of large-end design alternatives is diminished, and it
641 appears that that the high costs of investment in large design alternatives are not justified by their
642 relatively higher vulnerabilities to change.



643

644 Figure 8 (a) fraction scenarios in failure state (NPV<0) vs boxplot of all dry season energy
 645 production (average across all years of particular scenario); (b) maximum regret vs boxplot
 646 of all dry season energy production (average across all years of particular scenario).

647 Regret is the difference between the performance of some strategy in a particular future
 648 and the performance of the best strategy in that future (Savage, 1954). The design alternative that
 649 minimizes the maximum NPV regret across all tested combinations of climate conditions is the
 650 1000 MW option (Figure 8(b)). The 2000 MW design has potential in the very wet futures (right-
 651 side whiskers and outliers of the boxplots in Figure 8(b)) to produce much more hydropower,
 652 and consequently generate much greater NPV than the 335 MW design, but the median NPV for
 653 the 2000 MW design is negative across the modeled uncertainty range, strongly discouraging its
 654 adoption. Interestingly, maximum regret for the two extreme designs is similar, with greatest
 655 regrets for the 335 MW design occurring in wet futures, and greatest regret for the 2000 MW
 656 design occurring in dry futures. The 1355 MW design performs similarly to the 1000 MW in

657 terms of dry season hydropower production and maximum NPV regret, but requires 35 percent
658 more up-front investment capital, resulting in a preference for the 1000 MW design.

659 In the final comparison, the co-occurrence of the three conditions to which the 1000 MW
660 design is vulnerable (Table 3) is more likely than the co-occurrence of the three conditions to
661 which the 335 MW design option is vulnerable. According to the ensemble of GCM projections,
662 there is reason to believe that future precipitation might be less than 110% of historic, entering
663 the vulnerability range for the 1000 MW design. It is also reasonable to imagine that wet season
664 electricity price could be less than 0.10 US\$/kWh (more than double what it is now), and that
665 construction delays might result in cost overruns of 80% (on average, hydropower investments in
666 Nepal exceed their budgets by 60%). The preferability of the 1000 MW facility to the baseline
667 335 MW facility is contingent on a higher wet season selling price of electricity; investment in a
668 1000 MW hydropower facility is therefore something of a gamble on the ability of the India and
669 Nepal to come to terms. It is left to the investors and decision makers to determine whether the
670 opportunities presented by the 1000 MW design alternative outweigh the risks.

671 Other sets of conditions would threaten the performance of each design, and they should
672 not be ignored. For the 335 MW design, for example, though 82% of cases in which all three of
673 these conditions co-occur carry negative NPV (density), the co-occurrence of these three
674 conditions captures only 29% of the cases in which the project's NPV is negative (coverage).
675 Project lifetime, discount rate, plant load factor, and temperature change are also relevant, but in
676 the case of this particular investment are less important predictors for economic performance of
677 UAHP than capital cost, electricity price, and precipitation amount. For the 335 MW, 750 MW,
678 and 1000 MW options, the low coverage shown in Table 3 (in combination with relatively small
679 support) indicates that significant alternative failure scenarios exist, and should be explored in a

680 thorough treatment of system vulnerability, as explained in Annex 4 of Bonzanigo, et al. (2015),
681 available at [http://documents.worldbank.org/curated/en/179901476791918856/South-Asia-](http://documents.worldbank.org/curated/en/179901476791918856/South-Asia-Investment-decision-making-in-hydropower-decision-tree-case-study-of-the-upper-Arun-hydropower-project-and-Koshi-basin-hydropower-development-in-Nepal)
682 [Investment-decision-making-in-hydropower-decision-tree-case-study-of-the-upper-Arun-](http://documents.worldbank.org/curated/en/179901476791918856/South-Asia-Investment-decision-making-in-hydropower-decision-tree-case-study-of-the-upper-Arun-hydropower-project-and-Koshi-basin-hydropower-development-in-Nepal)
683 [hydropower-project-and-Koshi-basin-hydropower-development-in-Nepal](http://documents.worldbank.org/curated/en/179901476791918856/South-Asia-Investment-decision-making-in-hydropower-decision-tree-case-study-of-the-upper-Arun-hydropower-project-and-Koshi-basin-hydropower-development-in-Nepal).

684 **3.3. Discussion**

685 Throughout the course of the project evaluation, two design alternatives outperformed the
686 others, and the choice between those two design alternatives is subject to the risk/reward
687 preferences of the investors and stakeholders. The 1000 MW facility would produce more dry
688 season hydropower (and substantially more wet season hydropower) than the 335 MW facility;
689 however, it is more vulnerable to capital cost overruns, a weak electricity selling price, and the
690 possibility that the future might be drier than the past. As the median design alternative, the
691 maximum regret of the 1000 MW design is low, as its moderate size hedges against the
692 possibility of both a wetter and drier future. The 335 MW facility is a less ambitious capital
693 investment, and would be resilient to changes of all kinds, but it forfeits the potential to
694 capitalize on a potentially greater future streamflow condition (hence the high potential
695 maximum regret in the event of a wet future).

696 The case study is limited in its framing of the vulnerability-specific scenarios. Evaluation
697 of neither of the preferred designs (1000 MW or 335 MW) resulted in simple vulnerability-
698 specific scenarios with high coverage and density. Other parameters are needed to explain the
699 future conditions under which the project may fail, and other scenarios are needed to explain the
700 vulnerable futures. In some cases, no single scenario can provide adequate coverage and density.
701 In such cases, it may be necessary to iterate by identifying a vulnerability-specific scenario of
702 concern, removing overlapping cases, and rerunning the data mining algorithm to identify other

703 scenarios from the remaining set. The resulting set of multiple scenarios may reduce
704 interpretability, but can increase coverage and density. The end result is that the vulnerability-
705 specific scenarios are not equally relevant to the designs.

706 As a further limitation in characterizing risks, this study has evaluated the impacts on
707 project performance of climate-change induced shifts in temperature and precipitation, and not
708 climate-change induced exacerbation of precipitation extremes (flood and drought). The focus on
709 mean annual temperature and precipitation shifts is responsive to the level of confidence in the
710 local historical precipitation record (National Research Council, 2012), the best use of the
711 available climate projection science (Brown and Wilby, 2012), and the lack of skill of current
712 models (and associated downscaling techniques) in reproducing precipitation extremes (Cannon,
713 et al., 2015). However, this study may conclude that climate is of relatively low significance
714 because changes in the intensity and frequency of wet periods and dry periods have not been
715 adequately explored. Improvements to the current work would therefore incorporate risks of this
716 type, though it is unclear whether the available climate science would yet support such an
717 analysis.

718 Expert judgment in combination with analysis of climate trends and projections from the
719 latest ensemble of IPCC climate change scenarios can be used to assess the degree to which the
720 scenarios shown in Table 3 present a concern. Level of concern discussions beyond what has
721 been presented with reference to climate informed decision analysis, while important, are outside
722 of the scope of this analysis, though the results indicate that it would be very helpful were it
723 possible to place likelihoods on future conditions, despite the existence of irreducible
724 uncertainties that cannot easily be described by probability distributions. Can it reasonably be
725 said that a tripling of capital costs is equally likely to a doubling of capital costs? Including a

726 structured approach to uncertainty propagation, such as Bayesian Belief Networks, can inform
727 probabilistic weighting of the various change effects, and therefore a more useful description of
728 investment opportunities and risks across a wide range of possible realizations of an uncertain
729 future.

730 We note that there are a number of important considerations related to hydropower that
731 we have not adequately considered in this case study, although these are more political or
732 otherwise do not lend themselves to quantitative analysis: political disputes on ownership of
733 transboundary river waters (Grey and Sadoff, 2007; Grumbine, et al., 2012; He Daming, et al.,
734 2006; Kattelus, et al., 2015; Yoffe, et al., 2003); evidence of environmental damage (flooding of
735 animal habitat, obstruction of fish migration, and interference with natural stream ebbs and
736 flows) (Abbasi and Abbasi, 2011; Richter, et al., 1997; Schilt, 2007); and social costs of forced
737 displacement of valley inhabitants (e.g., Brown and Xu, 2010; IHA, 2016). Nonetheless they
738 require resolution in many cases to ensure that hydropower investments produce equitable
739 outcomes. Were these factors properly evaluated, and value given to the perspective of electricity
740 consumers instead of only electricity producers, the financial model presented in this work would
741 qualify as an economic model.

742 **4. CONCLUSIONS**

743 This paper has presented a generic methodology for evaluating climate change risks to
744 hydropower investments that simultaneously evaluates risks of many types in a multi-
745 dimensional stress test. The methodology helps to achieve a fundamental understanding of
746 climate changes on the response of a hydropower design.

747 The associated application to a proposed hydropower facility on the Arun River in Nepal
748 has demonstrated a direct comparison of alternative hydropower designs in response to the

749 concerns of relevance to the investors and stakeholders, and in answer to the call of Mukheibir
750 (2013), Gaudard and Romerio (2014), and others, for projected changes in the hydrological
751 regimes (and other relevant uncertainties) to be accounted for in future hydropower management
752 strategies and plans. The application illustrated an example in which climate change is not the
753 critical future uncertainty, and consequently highlighted the importance of considering multiple
754 uncertainties in combination. Finally, the climate change stress test undergirding this
755 multidimensional stress test accounted for the climate-change-induced shift in peak meltwater
756 contribution from April-May to March-April, sharpening the preference for designs not overly-
757 leveraged toward wet season hydropower productivity.

758 A note is warranted on the use of the term “resilient” in this work. In the general case,
759 hydropower projects can be evaluated for a variety of performance metrics, including robustness,
760 sustainability, flexibility, recovery, and others. The general framework presented in this paper
761 has adopted the definition of resilience put forth by the IHA, which includes both robustness and
762 recovery, though the case study has evaluated only the robustness aspects of resilience. Future
763 work applying this general framework will evaluate a wider set of perspectives on resilience.

764 Further research is also needed in the cryosphere – better monitoring – to inform
765 estimates of the available glacier mass, as well as how fast it is receding. Further research into
766 the sediment effects of precipitation extremes and the risks to the structure of natural hazards
767 (e.g., glacier lake outburst floods or landslides) is needed to better understand the changing
768 likelihood of such events with climate change.

769 Despite its limitations, the methodology presented in this paper has been useful to the
770 government of Nepal, which is using the results of the risk/reward tradeoff to inform decisions
771 about development of hydropower resources in the Koshi Basin.

772 **5. REFERENCES**

- 773 Abbasi T., Abbasi S. A. (2011) Small hydro and the environmental implications of its extensive
774 utilization. *Renewable & Sustainable Energy Reviews* 15, 2134-2143.
- 775 Akhtar M., Ahmad N., Booij M. J. (2008) The impact of climate change on the water resources
776 of Hindukush-Karakorum-Himalaya region under different glacier coverage scenarios.
777 *Journal of Hydrology* 355, 148-163.
- 778 Annandale G. W. (1987) Reservoir Sedimentation. 29, 221.
- 779 Ansar A., Flyvbjerg B., Budzier A., Lunn D. (2014) Should we build more large dams? The
780 actual costs of hydropower megaproject development. *Energy Policy* 69, 43-56.
- 781 Arrow K., Dasgupta P., Goulder L., Daily G., Ehrlich P., Heal G., Levin S., Maler K. G.,
782 Schneider S., Starrett D., Walker B. (2004) Are we consuming too much? *Journal of*
783 *Economic Perspectives* 18, 147-172.
- 784 Bajracharya S. R., Shrestha B. R. (2011) The Status of Glaciers in the Hindu Kush-Himalayan
785 Region. , 127.
- 786 Bankes S. (1993) Exploratory Modeling for Policy Analysis. *Operations Research* 41, 435-449.
- 787 Ben-Haim Y. (2006) Info-Gap Decision Theory: Decisions Under Severe Uncertainty. Academic
788 Press, London, UK.
- 789 Benn D. I., Bolch T., Hands K., Gulley J., Luckman A., Nicholson L. I., Quincey D., Thompson
790 S., Toumi R., Wiseman S. (2012) Response of debris-covered glaciers in the Mount
791 Everest region to recent warming, and implications for outburst flood hazards. *Earth-*
792 *Science Reviews* 114, 156-174.
- 793 Beyene T., Lettenmaier D. P., Kabat P. (2010) Hydrologic impacts of climate change on the Nile
794 River Basin: implications of the 2007 IPCC scenarios. *Climatic Change* 100, 433-461.

795 Bharati L., Gurung P., Jayakody P., Smakhtin V., Bhattarai U. (2014) The Projected Impact of
796 Climate Change on Water Availability and Development in the Koshi Basin, Nepal.
797 *Mountain Research and Development* 34, 118-130.

798 Bolch T., Kulkarni A., Kaab A., Huggel C., Paul F., Cogley J. G., Frey H., Kargel J. S., Fujita
799 K., Scheel M., Bajracharya S., Stoffel M. (2012) The State and Fate of Himalayan
800 Glaciers. *Science* 336, 310-314.

801 Bonzanigo L., Brown C., Harou J., Hurford A., Ray P., Karki P. (2015) South Asia investment
802 decision making in hydropower: Decision tree case study of the Upper Arun Hydropower
803 Project and Koshi Basin Hydropower Development in Nepal. *GEEDR South Asia, the*
804 *World Bank* Report No.: AUS 11077, 1-127.

805 Bookhagen B., Burbank D. W. (2010) Toward a complete Himalayan hydrological budget:
806 Spatiotemporal distribution of snowmelt and rainfall and their impact on river discharge.
807 *Journal of Geophysical Research-Earth Surface* 115, F03019.

808 Brown C., Wilby R. L. (2012) An alternate approach to assessing climate risks. *EOS,*
809 *Transactions, American Geophysical Union* 92, 401-412.

810 Brown C., Ghile Y., Lavery M., Li K. (2012) Decision scaling: Linking bottom-up vulnerability
811 analysis with climate projections in the water sector. *Water Resources Research* 48,
812 W09537.

813 Brown P. H., Xu K. (2010) Hydropower Development and Resettlement Policy on China's Nu
814 River. *Journal of Contemporary China* 19, 777-797.

815 Bush S. J., Turner A. G., Woolnough S. J., Martin G. M., Klingaman N. P. (2015) The effect of
816 increased convective entrainment on Asian monsoon biases in the MetUM general
817 circulation model. *Quarterly Journal of the Royal Meteorological Society* 141, 311-326.

818 Cannon A. J., Sobie S. R., Murdock T. Q. (2015) Bias Correction of GCM Precipitation by
819 Quantile Mapping: How Well Do Methods Preserve Changes in Quantiles and Extremes?
820 *Journal of Climate* 28, 6938-6959.

821 Castillo L. G., Carrillo J. M., Alvarez M. A. (2015) Complementary Methods for Determining
822 the Sedimentation and Flushing in a Reservoir. *Journal of Hydraulic Engineering* 141,
823 05015004.

824 Cervigni R., Liden R., Neumann J. E., Strzepek K. M. (2015) Enhancing the Climate Resilience
825 of Africa's Infrastructure: The Power and Water Sectors. , 1-192.

826 Chinnasamy P., Bharati L., Bhattarai U., Khadka A., Dahal V., Wahid S. (2015) Impact of
827 planned water resource development on current and future water demand in the Koshi
828 River basin, Nepal. *Water International* 40, 1004-1020.

829 Christensen N. S., Lettenmaier D. P. (2007) A multimodel ensemble approach to assessment of
830 climate change impacts on the hydrology and water resources of the Colorado River
831 Basin. *Hydrology and Earth System Sciences* 11, 1417-1434.

832 Christensen N. S., Wood A. W., Voisin N., Lettenmaier D. P., Palmer R. N. (2004) The effects
833 of climate change on the hydrology and water resources of the Colorado River basin.
834 *Climatic Change* 62, 337-363.

835 Dai F. C., Lee C. F., Deng J. H., Tham L. G. (2005) The 1786 earthquake-triggered landslide
836 dam and subsequent dam-break flood on the Dadu River, southwestern China.
837 *Geomorphology* 65, 205-221.

838 Davidson O., Metz B. (2000) IPCC Special Report - Emissions Scenarios: Summary for Policy
839 Makers. , 1-27.

840 Destouni G., Jaramillo F., Prieto C. (2013) Hydroclimatic shifts driven by human water use for
841 food and energy production. *Nature Climate Change* 3, 213-217.

842 Dussailant A., Benito G., Buytaert W., Carling P., Meier C., Espinoza F. (2010) Repeated
843 glacial-lake outburst floods in Patagonia: an increasing hazard? *Natural Hazards* 54, 469-
844 481.

845 Evans A., Strezov V., Evans T. J. (2009) Assessment of sustainability indicators for renewable
846 energy technologies. *Renewable & Sustainable Energy Reviews* 13, 1082-1088.

847 Finger D., Heinrich G., Gobiet A., Bauder A. (2012) Projections of future water resources and
848 their uncertainty in a glacierized catchment in the Swiss Alps and the subsequent effects
849 on hydropower production during the 21st century. *Water Resources Research* 48,
850 W02521.

851 Foster V., Briceño-Garmendia C. (2010) Africa's Infrastructure: A Time for Transformation. A
852 *Copublication of the Agence Française De Développement and the World Bank* , 1-40.

853 Friedman J. H., Fisher N. I. (1999) Bump hunting in high-dimensional data. *Statistics and*
854 *Computing* 9, 123-143.

855 Gaudard L., Romerio F. (2014) The future of hydropower in Europe: Interconnecting climate,
856 markets and policies. *Environmental Science & Policy* 37, 172-181.

857 Gaudard L., Gabbi J., Bauder A., Romerio F. (2016) Long-term Uncertainty of Hydropower
858 Revenue Due to Climate Change and Electricity Prices. *Water Resources Management*
859 30, 1325-1343.

860 Giuliani M., Anghileri D., Castelletti A., Phuong Nam Vu, Soncini-Sessa R. (2016) Large
861 storage operations under climate change: expanding uncertainties and evolving tradeoffs.
862 *Environmental Research Letters* 11, 035009.

863 Goulder L. H., Williams III R. C. (2012) The choice of discount rate for climate change policy
864 evaluation. *Climate Change Economics* 3, 1-18.

865 Grey D., Sadoff C. (2007) Sink or swim? Water security for growth and development. *Water*
866 *Policy* 9, 545-571.

867 Grinsted A. (2013) An estimate of global glacier volume. *Cryosphere* 7, 141-151.

868 Groves D. G., Lempert R. J. (2007) A new analytic method for finding policy-relevant scenarios.
869 *Global Environmental Change-Human and Policy Dimensions* 17, 73-85.

870 Groves D., Mao Z., Liden R., Strzepek K. M., Lempert R., Brown C., Taner M., Bloom E.
871 (2015) Adaptation to Climate Change in Project Design. , 131-154.

872 Groves D. G., Yates D., Tebaldi C. (2008) Developing and applying uncertain global climate
873 change projections for regional water management planning. *Water Resources Research*
874 44, W12413.

875 Grumbine R. E., Dore J., Xu J. (2012) Mekong hydropower: drivers of change and governance
876 challenges. *Frontiers in Ecology and the Environment* 10, 91-98.

877 Hamlet A. F., Lee S., Mickelson K. E. B., Elsner M. M. (2010) Effects of projected climate
878 change on energy supply and demand in the Pacific Northwest and Washington State.
879 *Climatic Change* 102, 103-128.

880 Hanshaw M. N., Bookhagen B. (2014) Glacial areas, lake areas, and snow lines from 1975 to
881 2012: status of the Cordillera Vilcanota, including the Quelccaya Ice Cap, northern
882 central Andes, Peru. *Cryosphere* 8, 359-376.

883 He Daming, Feng Yan, Gan Shu, Magee D., You Weihong. (2006) Transboundary hydrological
884 effects of hydropower dam construction on the Lancang River. *Chinese Science Bulletin*
885 51, 16-24.

886 Ho J. T., Thompson J. R., Brierley C. (2016) Projections of hydrology in the Tocantins-Araguaia
887 Basin, Brazil: uncertainty assessment using the CMIP5 ensemble. *Hydrological Sciences*
888 *Journal-Journal Des Sciences Hydrologiques* 61, 551-567.

889 Hosterman H. R., McCornick P. G., Kistin E. J., Sharma B., Bharati L. (2012) Freshwater,
890 climate change and adaptation in the Ganges River Basin. *Water Policy* 14, 67-79.

891 IDS-Nepal P., and GCAP. (2014) Economic Impact Assessment of Climate Change in Key
892 Sectors in Nepal: Summary Report. *Integrated Development Society Nepal (IDS-Nepal),*
893 *Practical Action Consulting (PAC), and the Global Climate Adaptation Partnership*
894 *(GCAP)* Apr 2014, 1-24.

895 IEG. (2012) Adapting to Climate Change: Assessing the World Bank Group Experience.

896 IHA. (2016) Hydropower Status Report. *International Hydropower Association* , 1-79.

897 Immerzeel W. W., Pellicciotti F., Bierkens M. F. P. (2013) Rising river flows throughout the
898 twenty-first century in two Himalayan glacierized watersheds. *Nature Geoscience* 6, 742-
899 745.

900 Immerzeel W. W., van Beek, Ludovicus P H, Bierkens M. F. P. (2010) Climate Change Will
901 Affect the Asian Water Towers. *Science* 328, 1382-1385.

902 Jaramillo O. A., Boria M. A., Huacuz J. M. (2004) Using hydropower to complement wind
903 energy: a hybrid system to provide firm power. *Renewable Energy* 29, 1887-1909.

904 Kaab A., Berthier E., Nuth C., Gardelle J., Arnaud Y. (2012) Contrasting patterns of early
905 twenty-first-century glacier mass change in the Himalayas. *Nature* 488, 495-498.

906 Kasprzyk J. R., Nataraj S., Reed P. M., Lempert R. J. (2013) Many objective robust decision
907 making for complex environmental systems undergoing change. *Environmental*
908 *Modelling & Software* 42, 55-71.

909 Kattelus M., Kummu M., Keskinen M., Salmivaara A., Varis O. (2015) China's southbound
910 transboundary river basins: a case of asymmetry. *Water International* 40, 113-138.

911 Kilsby C. G., Jones P. D., Burton A., Ford A. C., Fowler H. J., Harpham C., James P., Smith A.,
912 Wilby R. L. (2007) A daily weather generator for use in climate change studies.
913 *Environmental Modeling & Software* 22, 1705-1719.

914 Kubiszewski I., Costanza R., Paquet P., Halimi S. (2013) Hydropower development in the lower
915 Mekong basin: alternative approaches to deal with uncertainty. *Regional Environmental*
916 *Change* 13, 3-15.

917 Kucukali S. (2011) Risk assessment of river-type hydropower plants using fuzzy logic approach.
918 *Energy Policy* 39, 6683-6688.

919 Kwakkel J. H., Haasnoot M., Walker W. E. (2016) Comparing Robust Decision-Making and
920 Dynamic Adaptive Policy Pathways for model-based decision support under deep
921 uncertainty. *Environmental Modelling & Software* 86, 168-183.

922 Laghari A. N., Vanham D., Rauch W. (2012) To what extent does climate change result in a shift
923 in Alpine hydrology? A case study in the Austrian Alps. *Hydrological Sciences Journal-*
924 *Journal Des Sciences Hydrologiques* 57, 103-117.

925 Latham J., Cumani R., Rosati L., Bloise M. (2014) FAO Global Land Cover (GLC-SHARE).
926 *Land and Water Division Beta-Release 1.0 Database*, 1-40.

927 Lehner B., Czisch G., Vassolo S. (2005) The impact of global change on the hydropower
928 potential of Europe: a model-based analysis. *Energy Policy* 33, 839-855.

929 Lempert R. J., Groves D. G., Popper S. W., Bankes S. C. (2006) A general, analytic method for
930 generating robust strategies and narrative scenarios. *Management Science* 52, 514-528.

- 931 Lempert R. J., Popper S. W., Bankes S. C. (2003) *Shaping the Next One Hundred Years : New*
932 *Methods for Quantitative, Long-Term Policy Analysis*. RAND Corporation, Santa
933 Monica.
- 934 Loo Y. Y., Billa L., Singh A. (2015) Effect of climate change on seasonal monsoon in Asia and
935 its impact on the variability of monsoon rainfall in Southeast Asia. *Geoscience Frontiers*
936 6, 817-823.
- 937 Loucks D. P., Van Beek E. (2005) *Water Resources Systems Planning and Management: An*
938 *Introduction to Methods, Models and Applications*. UNESCO Publishing, Paris, France.
- 939 Lutz A. F., Immerzeel W. W., Shrestha A. B., Bierkens M. F. P. (2014) Consistent increase in
940 High Asia's runoff due to increasing glacier melt and precipitation. *Nature Climate*
941 *Change* 4, 587-592.
- 942 Madani K., Lund J. R. (2010) Estimated impacts of climate warming on California's high-
943 elevation hydropower. *Climatic Change* 102, 521-538.
- 944 Majone B., Villa F., Deidda R., Bellin A. (2016) Impact of climate change and water use policies
945 on hydropower potential in the south-eastern Alpine region. *Science of the Total*
946 *Environment* 543, 965-980.
- 947 Manning L. J., Hall J. W., Fowler H. J., Kilsby C. G., Tebaldi C. (2009) Using probabilistic
948 climate change information from a multimodel ensemble for water resources assessment.
949 *Water Resources Research* 45, W11411.
- 950 Markoff M. S., Cullen A. C. (2008) Impact of climate change on Pacific Northwest hydropower.
951 *Climatic Change* 87, 451-469.

952 Maurer E. P., Adam J. C., Wood A. W. (2009) Climate model based consensus on the hydrologic
953 impacts of climate change to the Rio Lempa basin of Central America. *Hydrology and*
954 *Earth System Sciences* 13, 183-194.

955 McKay M. D., Beckman R. J., Conover W. J. (1979) A Comparison of Three Methods for
956 Selecting Values of Input Variables in the Analysis of Output from a Computer Code.
957 *Technometrics* 21, 239-245.

958 Mehta V. K., Rheinheimer D. E., Yates D., Purkey D. R., Viers J. H., Young C. A., Mount J. F.
959 (2011) Potential impacts on hydrology and hydropower production under climate
960 warming of the Sierra Nevada. *Journal of Water and Climate Change* 2, 29-43.

961 Mendelsohn R. O. (2008) Is the Stern Review an economic analysis? *Review of Environmental*
962 *Economics and Policy* 2, 45-60.

963 Minville M., Brissette F., Krau S., Leconte R. (2009) Adaptation to Climate Change in the
964 Management of a Canadian Water-Resources System Exploited for Hydropower. *Water*
965 *Resources Management* 23, 2965-2986.

966 Minville M., Brissette F., Leconte R. (2008) Uncertainty of the impact of climate change on the
967 hydrology of a nordic watershed. *Journal of Hydrology* 358, 70-83.

968 Mishra S., Singal S. K., Khatod D. K. (2011) Optimal installation of small hydropower plant-A
969 review. *Renewable & Sustainable Energy Reviews* 15, 3862-3869.

970 Mukheibir P. (2013) Potential consequences of projected climate change impacts on
971 hydroelectricity generation. *Climatic Change* 121, 67-78.

972 National Research Council. (2012) Himalayan Glaciers: Climate change, water resources, and
973 water security
974 . *The National Academic Press, Washington, D C* , 1-142.

975 NEA. (2014) Brief Introduction on Nepal's Hydro Potential and Power Development.

976 NEA. (1991) Upper Arun Hydropower Project Feasibility Study. *Joint Venture of Morrison*
977 *Knudsen Corporation, Lahmeyer International, Tokyo Electric Power Services Co , and*
978 *NEPECON .*

979 Nepal S., Shrestha A. B. (2015) Impact of climate change on the hydrological regime of the
980 Indus, Ganges and Brahmaputra river basins: a review of the literature. *International*
981 *Journal of Water Resources Development* 31, 201-218.

982 Nordhaus W. (2007) A review of the Stern Review on the economics of climate change. *Journal*
983 *of Economic Literature* 45, 686-702.

984 OECD IEA. (2015) Key World Energy Statistics. *International Energy Agency* 1, 1-81.

985 Peng M., Zhang L. M. (2012) Analysis of human risks due to dam-break floods-part 1: a new
986 model based on Bayesian networks. *Natural Hazards* 64, 903-933.

987 Pfeffer W. T., Arendt A. A., Bliss A., Bolch T., Cogley J. G., Gardner A. S., Hagen J., Hock R.,
988 Kaser G., Kienholz C., Miles E. S., Moholdt G., Moelg N., Paul F., Radic V., Rastner P.,
989 Raup B. H., Rich J., Sharp M. J., Randolph Consortium. (2014) The Randolph Glacier
990 Inventory: a globally complete inventory of glaciers. *Journal of Glaciology* 60, 537-552.

991 Prudhomme C., Wilby R. L., Crooks S., Kay A. L., Reynard N. S. (2010) Scenario-neutral
992 approach to climate change impact studies: Application to flood risk. *Journal of*
993 *Hydrology* 390, 198-209.

994 Rasul G. (2014) Food, water, and energy security in South Asia: A nexus perspective from the
995 Hindu Kush Himalayan region. *Environmental Science & Policy* 39, 35-48.

996 Ray P. A., Brown C. M. (2015) Confronting Climate Uncertainty in Water Resources Planning
997 and Project Design: The Decision Tree Framework. World Bank, Washington, DC.

- 998 Ray P. A., Yang Y. E., Wi S., Khalil A., Chatikavanij V., Brown C. (2015) Room for
999 improvement: Hydroclimatic challenges to poverty-reducing development of the
1000 Brahmaputra River basin. *Environmental Science & Policy* 54, 64-80.
- 1001 Richardson S. D., Reynolds J. M. (2000) An overview of glacial hazards in the Himalayas.
1002 *Quaternary International* 65-6, 31-47.
- 1003 Richter B. D., Braun D. P., Mendelson M. A., Master L. L. (1997) Threats to imperiled
1004 freshwater fauna. *Conservation Biology* 11, 1081-1093.
- 1005 Rocheta E., Sugiyanto M., Johnson F., Evans J., Sharma A. (2014) How well do general
1006 circulation models represent low-frequency rainfall variability? *Water Resources*
1007 *Research* 50, 2108-2123.
- 1008 Rosenberg D. M., Bodaly R. A., Usher P. J. (1995) Environmental and Social Impacts of Large-
1009 Scale Hydroelectric Development - Who is Listening. *Global Environmental Change-
1010 Human and Policy Dimensions* 5, 127-148.
- 1011 Saltelli A., Chan K., Scott E. M. (2009) Sensitivity Analysis. Wiley, New York, USA.
- 1012 Santolin A., Cavazzini G., Pavesi G., Ardizzon G., Rossetti A. (2011) Techno-economical
1013 method for the capacity sizing of a small hydropower plant. *Energy Conversion and
1014 Management* 52, 2533-2541.
- 1015 Savage L. J. (1954) The Foundation of Statistics. Dover Publications.
- 1016 Savoskul O. S., Smakhtin V. (2013) Glacier Systems and Seasonal Snow Cover in Six Major
1017 Asian River Basins: Hydrological Role under Changing Climate. *International Water
1018 Management Institute Report* 150, IWMI, Colombo, Sri Lanka, 1-45.

1019 Schaefli B., Hingray B., Musy A. (2007) Climate change and hydropower production in the
1020 Swiss Alps: quantification of potential impacts and related modelling uncertainties.
1021 *Hydrology and Earth System Sciences* 11, 1191-1205.

1022 Schilt C. R. (2007) Developing fish passage and protection at hydropower dams. *Applied Animal*
1023 *Behaviour Science* 104, 295-325.

1024 Schneider U., Becker A., Finger P., Meyer-Christoffer A., Ziese M., Rudolf B. (2014) GPCC's
1025 new land surface precipitation climatology based on quality-controlled in situ data and its
1026 role in quantifying the global water cycle. *Theoretical and Applied Climatology* 115, 15-
1027 40.

1028 Sharma R. H., Shakya N. M. (2006) Hydrological changes and its impact on water resources of
1029 Bagmati watershed, Nepal. *Journal of Hydrology* 327, 315-322.

1030 Shrestha A. B., Aryal R. (2011) Climate change in Nepal and its impact on Himalayan glaciers.
1031 *Regional Environmental Change* 11, S77.

1032 Singh P., Bengtsson L. (2004) Hydrological sensitivity of a large Himalayan basin to climate
1033 change. *Hydrological Processes* 18, 2363-2385.

1034 Sorg A., Bolch T., Stoffel M., Solomina O., Beniston M. (2012) Climate change impacts on
1035 glaciers and runoff in Tien Shan (Central Asia). *Nature Climate Change* 2, 725-731.

1036 Stainforth D. A., Downing T. E., Washington R., Lopez A., New M. (2007) Issues in the
1037 interpretation of climate model ensembles to inform decisions. *Philosophical*
1038 *Transactions of the Royal Society A-Mathematical Physical and Engineering Sciences*
1039 365, 2163-2177.

1040 Steinschneider S., Brown C. (2013) A semiparametric multivariate, multisite weather generator
1041 with low-frequency variability for use in climate risk assessments. *Water Resources*
1042 *Research* 49, 7205-7220.

1043 Steinschneider S., McCrary R., Mearns L. O., Brown C. (2015) The effects of climate model
1044 similarity on probabilistic climate projections and the implications for local, risk-based
1045 adaptation planning. *Geophysical Research Letters* 42, 5014-5022.

1046 Stouffer R. J., Eyring V., Meehl G. A., Bony S., Senior C., Stevens B., Taylor K. E. (2017)
1047 Cmp5 Scientific Gaps and Recommendations for Cmp6. *Bulletin of the American*
1048 *Meteorological Society* 98, +.

1049 Tan P., Steinbach M., Kumar V. (2005) Introduction to Data Mining. Pearson Addison Wesley,
1050 Boston.

1051 Taylor K. E., Stouffer R. J., Meehl G. A. (2012) An Overview of Cmp5 and the Experiment
1052 Design. *Bulletin of the American Meteorological Society* 93, 485-498.

1053 Tebaldi C., Smith R. L., Nychka D., Mearns L. O. (2005) Quantifying uncertainty in projections
1054 of regional climate change: A Bayesian approach to the analysis of multimodel
1055 ensembles. *Journal of Climate* 18, 1524-1540.

1056 Turner A. G., Annamalai H. (2012) Climate change and the South Asian summer monsoon.
1057 *Nature Climate Change* 2, 587-595.

1058 van Vliet, Michelle T H, Wiberg D., Leduc S., Riahi K. (2016) Power-generation system
1059 vulnerability and adaptation to changes in climate and water resources. *Nature Climate*
1060 *Change* 6, +.

1061 Vogel R. M., Shallcross A. L. (1996) The moving blocks bootstrap versus parametric time series
1062 models. *Water Resources Research* 32, 1875-1882.

1063 Wang H., Long L., Kumar A., Wang W., Schemm J. E., Zhao M., Vecchi G. A., Larow T. E.,
1064 Lim Y., Schubert S. D., Shaevitz D. A., Camargo S. J., Henderson N., Kim D., Jonas J.
1065 A., Walsh K. J. E. (2014) How Well Do Global Climate Models Simulate the Variability
1066 of Atlantic Tropical Cyclones Associated with ENSO? *Journal of Climate* 27, 5673-
1067 5692.

1068 Wi S., Yang Y. C. E., Steinschneider S., Khalil A. F., Brown C. M. (2015) Calibration
1069 approaches for distributed hydrologic models in poorly gaged basins: Implication for
1070 streamflow projections under climate change. *Hydrology and Earth System Sciences* 19,
1071 857-876.

1072 Wilby R. L., Dawson C. W. (2013) The Statistical DownScaling Model: insights from one
1073 decade of application. *International Journal of Climatology* 33, 1707-1719.

1074 Wilby R. L., Harris I. (2006) A framework for assessing uncertainties in climate change impacts:
1075 Low-flow scenarios for the River Thames, UK. *Water Resources Research* 42, W02419.

1076 Wild T. B., Loucks D. P., Annandale G. W., Kaini P. (2016) Maintaining Sediment Flows
1077 through Hydropower Dams in the Mekong River Basin. *Journal of Water Resources*
1078 *Planning and Management* 142, 05015004.

1079 Wilks D. S., Wilby R. L. (1999) The weather generation game: a review of stochastic weather
1080 models. *Progress in Physical Geography* 23, 329-357.

1081 World Bank T. (2016) Discounting costs and benefits in economic analysis of World Bank
1082 projects. *Operations Policy and Quality Department (OPSPQ)* 9 May, 1-10.

1083 World Bank T. (2009) Directions in hydropower. *Technical Report* .

1084 World Energy Council. (2016) World Energy Resources: Hydropower 2016. *World Energy*
1085 *Council* , 1-53.

1086 World Energy Council. (2013) World Energy Resources: 2013 Survey. *World Energy Council* ,
1087 1-468.

1088 Yang Y. C. E., Wi S., Ray P. A., Brown C. M., Khalil A. F. (2016) The future nexus of the
1089 Brahmaputra River Basin: Climate, water, energy and food trajectories. *Global*
1090 *Environmental Change-Human and Policy Dimensions* 37, 16-30.

1091 Yatagai A., Kamiguchi K., Arakawa O., Hamada A., Yasutomi N., Kitoh A. (2012)
1092 APHRODITE: Constructing a Long-Term Daily Gridded Precipitation Dataset for Asia
1093 Based on a Dense Network of Rain Gauges. *Bulletin of the American Meteorological*
1094 *Society* 93, 1401-1415.

1095 Yates D., Gangopadhyay S., Rajagopalan B., Strzepek K. (2003) A technique for generating
1096 regional climate scenarios using a nearest-neighbor algorithm. *Water Resources Research*
1097 39, 1199.

1098 Yoffe S., Wolf A. T., Giordano M. (2003) Conflict and cooperation over international freshwater
1099 resources: Indicators of basins at risk. *Journal of the American Water Resources*
1100 *Association* 39, 1109-1126.

1101 Zakeri B., Syri S. (2015) Electrical energy storage systems: A comparative life cycle cost
1102 analysis. *Renewable & Sustainable Energy Reviews* 42, 569-596.

1103 Zarfl C., Lumsdon A. E., Berlekamp J., Tydecks L., Tockner K. (2015) A global boom in
1104 hydropower dam construction. *Aquatic Sciences* 77, 161-170.

1105 Zhuang J., Liang Z., Lin T., De Guzman F. (2007) Theory and Practice in the Choice of Social
1106 Discount Rate for Cost-Benefit Analysis: A Survey. *Economics and Research*
1107 *Department Working Paper, Asian Development Bank* 94, 1-40.

1108